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5 **Stability evaluation of streambanks in Vermont**

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25 **Abstract**

26 Streambank stability characterization has become a topic of major interest with rising
27 concerns regarding phosphorous accumulation in surface water bodies. The ability to make rapid
28 reliable determination of a streambank's stability is necessary as the ability to make informed
29 land management decisions becomes paramount. Since the summer of 2006 two stream reaches
30 in the Lake Champlain basin of northern Vermont have been characterized and monitored
31 through the use of surveying, geotechnical investigations, and instrumentation. Temporal
32 survey results were used to observe mass failures, sediment removal and deposition along the
33 streambanks throughout the course of this study. Laboratory testing was used to determine soils,
34 strengths, root strengths and soil erosion characteristics. Stability of the streambanks was
35 analyzed using classical soil mechanics incorporating variable stream and ground water tables
36 and results from the laboratory testing; revealing year round instability of many of the
37 streambanks observed. This study showed the validity of the Iowa borehole shear test and the
38 necessity to incorporate matric suction additions to soils strength.

39

40 **1. Introduction**

41 A large portion of potable water is contained in the small percentage of fresh surface
42 water found on the surface of the earth. Today these fresh water sources are under constant
43 threat from accelerated pollution due to anthropocentric causes. Pollution found in surface waters
44 is caused by toxic and conventional pollutants (Kreger 2004). Although toxins pose immediate
45 and serious health threats represent only a small portion of the total pollution found in our
46 waterway. Conventional pollution represents the larger portion of pollution found today, being
47 defined as nutrients and biological contaminants that enter waterways through natural means. Of
48 pollutants phosphorus (P) has become recognized as the largest in US lakes and water
49 ways(Agency 2009). Anthropocentric activity has been shown to increase conventional
50 pollution levels; particularly in levels of nitrogen and phosphorous (Carpenter 1998).
51 Responsible for these higher loading rates of conventional pollutants are commonly sewage

52 treatment effluent, stormwater runoff, and non-point sources associated with agricultural and
53 land management practices.

54 Large inorganic sources today are both point and non-point agricultural sources of
55 particular importance due to increasing fertilizer application rates (Holtan 1988). Due to
56 phosphorus nature to tightly bind to soil particles(Barros 2005); P transport is closely linked with
57 that alluvial sediments(cite). Particulate-bound phosphorous (PP) levels quantified at the
58 outflow of a watershed is a combination of PP that has entered into the waterways recently via
59 point and non point sources, remobilized PP from alluvial deposits through stream bed
60 migration and erosion processes. Streambank erosion including scour and mass failure is
61 estimated to account for 30-80% of sediment loading into lakes and waterways (1999; Evans
62 2006; Fox 2007).

63 Much is known about sediment source processes, transport, and deposition. However
64 modeling of these processes has applicability only to study regions where sources of sediments
65 may be easily identified. Several direct measurement procedures have been used in efforts to
66 quantify bank erosion. Lawler (1999) made use of longitudinal surveys and pins to quantify
67 sediment loading though bank erosion. Longitudinal surveys allowed the measurement of bank
68 top retreats while the use of pins allowed measurement of toe erosion of laterally migrating
69 streambanks. As such direct techniques have been found valuable in determining sediment loads
70 from sources in small watersheds that can usually be readily recognized(cite). Labor intensity of
71 direct measurement techniques makes large scale use infeasible. Due to the labor of purely
72 quantitative methods and waning confidence in qualitative methods it is necessary to develop a
73 semi-quantitative approach to determine bank stability. Such an approach would make use of
74 bulk samples and inexpensive testing procedures to determine soil types which would allow for

75 determination of material properties of soil through the use of empirical calculations and index
76 properties.

77 The following study was preformed as a initial efforts on developing a method for
78 determining streambank stabilities in the alluvial river beds of Vermont. A series of surveys were
79 conducted to monitor geometric changes in river cut banks. Subsurface investigations were also
80 preformed in efforts to characterized local soil deposits type, and strength. Data gathered from
81 these studies were used to analyze the changing stability of river cut banks using classical soil
82 mechanics.

83

84 **2. Factors Affecting Streambank Stability**

85 2.1 Geotechnical Material strength

86 The most visible process that causes streambank changes is that of mass failure. These
87 failures can be analyzed by slope stability methods. Several methods of analyzing slope stability
88 are applicable to assessing the stability of streambanks for a single point in time (Langedoen et
89 al., 2001). Bishop (1955) developed a technique to assess slope stability on a circular failure
90 surface. Although failure surfaces are often circular non-circular failure surfaces occur more
91 commonly in streambanks where common failure geometries are slab, rotational, wedge, and
92 cantilever with a dependence on the amount of bank undercutting prior to failure(Simon and
93 Collision 2002). To account for non-circular failure surfaces an analysis technique was
94 developed that divides a predefined failure surface into a number of finite slices, where the
95 driving and resisting forces along each section of the slip surface for each slice are quantified and

96 summed (Morgenstern and Price 1965). Values from this method are then looked at in terms of
97 a factor of safety; defined as the ratio of resisting forces to the driving forces. To quantify the
98 resisting forces of the soil the Mohr-Coulomb failure criterion is used (equation 1).

$$99 \quad \tau = c' + (\sigma_n - \mu_w) \tan \phi' \quad (1)$$

100 Where c' is the effective cohesion of the soil; σ_n is the confining pressure on the slip surface; μ_w
101 is pore water pressure, and ϕ' is the soils effective angle of internal friction. If the driving forces
102 exceed the resisting forces for the given failure surface a factor of safety of less than one is seen
103 indicating slope instability and failure. These methods tend to provide accurate analysis for
104 instantaneous slope stability of a given geometry. However, this system is used to predict the
105 stability of saturated slopes where only non-negative pore water pressures are allowed to
106 develop.

107

108 2.2 Vegetation effects

109 Slope stability of streambanks is inherently more difficult to calculate than that of an
110 earthen structure. Another aspect that needs to be considered when evaluating slope stability is
111 the effects of vegetation found on streambanks. Vegetation has two primary effects on the
112 stability of streambanks the surcharge imposed by vegetation on the bank and the shear strength
113 addition caused by the root structure in the soil (De Wiel and Darby 2007). The surcharge can
114 be added into a slope stability analysis quite easily by the addition of a pressure to the bank top.
115 However the strength addition caused by roots is more difficult to incorporate. It has long been
116 assumed that the presence of roots in soils increases the shear strength of soils, and therefore the
117 stability of streambanks(Collision 2002). The soil root system creates a composite material

118 where roots add tensile strength to the material while the soil provides compressive strength.
119 The added strength of the soil-root composite material is dependent upon the fractional volume
120 of live roots, tensile strength, and orientation of the roots. (Oshashi and Gray,1983; Wu *et al.*,
121 1979). The directional orientation of a root system through a failure surface has been shown to
122 change the strain at which maximum shear strength is encountered (Wu, McKinnell III et al.
123 1979; Ennos 1990). Furthermore the spatial distribution of roots, generally quantified as a root-
124 area ratio, is non uniform. With grass and sages that are densely grouped the root area ratio may
125 be assumed as linear function with a maximum root area ratio at the surface and degrading to
126 zero at the species maximum depth of rooting. With shrubs and trees the distribution is more
127 sporadic and the root area ratios decrease in both the horizontal and vertical directions from the
128 stem/trunk forming a cone shape root area ratio distribution that extend far beneath the surface
129 (De Wiel and Darby 2007).

130 Equations developed by (Waldron 1977) estimate strength addition to soils due to the
131 tensile properties of roots. Strength additions to a soil may be added as an increase in the
132 cohesion of the soil as modeled by the Mohr-Coulomb failure criterion (equation 2).

$$133 \quad \tau = c_r + c'_s + (\sigma_n - \mu_w) \tan \phi' \quad (2)$$

134 Where τ is the soil's shear strength; c_r is apparent cohesion due to roots; c'_s is the soil's effective
135 cohesion; σ_n is the effective confining pressure; μ_w is the pore-water pressure and ϕ' is the
136 effective internal shear angle of the soil. Waldron (1977) developed a generalized equation to
137 describe additional cohesion provided by roots (equation 3.). This equation was later simplified
138 after determining that for most variations of ϕ and θ a coefficient of 1.2 can be used with less
139 than 10% additional error(Wu, McKinnell III et al. 1979). As determination of the root area ratio

140 of soils leads to significant variability in estimated shear strength additions due to root systems
141 this simplification may be used with negligible effect on the results.

$$142 \quad c'_r = (\sin\theta + \cos\theta \tan\phi)(T_r R_a) \quad (3)$$

143 Where T_r is the tensile strength of roots; R_a is the root area ratio defined as the ratio of root
144 volume to total volume of a given soil sample; and θ is the angle of roots passing through the
145 failure plane.

146

147 2.3 Soil Pore-water Pressures

148 Pore-water pressures have been long known to affect the stability of an engineered slope,
149 with positive pore-water hydrostatic forces creating positive pressures below the phreatic
150 surface, with saturated conditions accounted for in the Mohr-Coulomb failure criterion (equation
151 1.). Due to the transient nature of stream water levels streambanks are often comprised of
152 regions of saturated and unsaturated soils. Above the phreatic surface of the water negative pore
153 water pressures develop which could affect the strength of the materials significantly and
154 therefore the stability of the slope. The effects of negative pore-water pressures (matric suction)
155 most easily are observed by the steep bank angles that occur at eroding stream sites. Matric
156 suction often develops during the lowering of water levels that causing disjuncting of negatively
157 pressurized water from the capillary fringe, and may be increased through effects of
158 dehydration and evapotranspiration(cite). Predictions of the magnitude of matric suction can
159 be made through use of the soil water characteristic curve (figure 2.1.), which relates the value of
160 matric suction to the degree of saturation of the soil.

161 For soils at 85% saturation or higher, at or below the air entry pressure, low values of
162 matric suction and uniform pore-water distribution provide for easy analysis. Soil strengths in
163 this zone can be analyzed using classical soil mechanics by using proper sign convention to
164 account the negative water pressures (Fredlund and Rahardjo 1987). For values of matric
165 suction greater than the air entry pressure the soil desaturates and pressures become
166 heterogeneous (Mohamed, Ali et al. 2006). Changes in water content and pressures at the
167 granular scale in the unsaturated zone were cause for a revision of the Mohr-Coulomb failure
168 criterion by Fredlund and Rahardjo(1987) to the form of equation 4 to account for these
169 inconsistencies.

$$170 \quad \tau = c' + (\sigma - \mu_a)\tan\phi' + \psi\tan\phi^b \quad (4)$$

171 Where τ is the unsaturated shear strength of the soil; c' is the effective cohesion of the soil; σ is
172 the stress applied to the soil; μ_a is the barometric pressure; ϕ' is the effective friction angle of the
173 soil; ψ is the matric suction pressure, and ϕ^b is the angle of strength change with variations in
174 matric suction.

175

176 2.4 Hydraulic effects

177 Another difficulty with predicting the long term stability of a stream bank is its constantly
178 changing shape. Most bank failures occur due to undercutting, with the failure mode geometry
179 strongly related to the degree of undercutting of the bank (Simon and Andrew 2001). Unlike
180 more noticeable changes in the stream this undercutting is not caused by mass failure, but by
181 scour erosion. The rate at which this erosion takes place is illustrated by the following equation.

182
$$E_r = k_d(\tau_e - \tau_c) \quad (5)$$

183 Where E_r is the erosion rate, k_d is the erosion coefficient, τ_e is the effective shear stress, and τ_c
184 is the soils critical shear stress (Hanson, 199#). Scour occurs when effective shear stress exerted
185 on the exposed portion of soil is greater than the critical shear stress of the soil causing soil
186 detachment (Hanson 199#). Empirical correlations for soils with a mean particle diameter
187 greater than 0.3mm the critical shear stress (pa) is equal to the mean grain size D_{50} (mm). For
188 soils with $D_{50}<0.3$ mm this equation is not useful for an accurate determination of critical shear
189 stress (Briaud J.L. 2001). In soils consisting of finer grains delicate intricacies make
190 predication of the erosion characteristics very difficult. Erosion of soils of this size is often
191 affected by difficult to quantify electro-chemical interactions that may create multi partial
192 erosion characteristics (Mazurek, Rajaratnam et al. 2001). In order to obtain critical shear stress
193 and erosion rate values direct experimental testing is needed.

194

195 **3. Site Selection and Instrumentation**

196 3.1 Site Selection

197 Several requirements were needed for a stream in this study. Due to the time necessary for
198 site characterization a multitude of site visits would need to be preformed. To accommodate this
199 only stream reaches within 50km (35miles) of Burlington , VT would be considered for this
200 study. Due to the nature of geotechnical investigations permission from landowners was also
201 necessary, which was conditionally obtained for an initial non-intrusive site visit. Access was
202 also a consideration as remote areas would prevent use of heavy drilling equipment. Reaches

203 from a previous study conducted by Hession et al. (year) meeting this criteria were individually
204 visited. Of the reaches that were investigated permission was granted to install instrumentation
205 at a number of these sites; however permission was not given for any heavy drilling equipment.

206 Of the reaches where instrumentation privileges were granted further selection was
207 conducted. Reaches with prevalent failing streambanks and streambanks that were marginally
208 stable were further considered. The height of the falling banks was also taken into consideration
209 with an optimal range of 2-5m (6.6-16.4ft) in height to provide sufficient survey resolution and
210 subsurface investigation using hand auguring equipment. The soil deposits of the area had to be
211 primarily composed of sand, silt, or clay. This provision was to insure that hand operated
212 drilling equipment and insitu strength tests could be readily conducted for a more accurate
213 quantification of the subsurface soils. Two sites were selected meeting these requirements. The
214 first site selected was a lower reach of the Winooski River located in Burlington, VT. The
215 second site is located in the lower reaches of Lewis Creek located in Ferrisburg, VT adjacent to
216 Lake Champlain (see figure 3.1.).

217 Seven cross sections were selected at each reach for temporal monitoring. In order to observe
218 changes in the streambank geometry a series of topographic surveys were used. Although time
219 consuming this method was selected for use instead of the less labor intensive pin method
220 developed by Lawler (year). This method allowed for changes in the entire cross section to be
221 noted which would be particularly important in characterizing slope failure geometries. The
222 population of cross sections observed was composed of streambanks that had recently failed or
223 were marginal stable. To facilitate multiple repeated surveys of the same cross sections pairs of
224 iron rods were placed perpendicular to the channel flow allowing for repeated surveys of the
225 same stream cross section. Survey sets were conducted using a Topcon GS 236 total station

226 during the summer of 2007 and 2008 with additional surveys during the fall of 2008. In one
227 instance a large bank failure prevented repetitive surveys of a cross section at the Lewis Creek
228 site as both survey pins were removed.

229

230 3.2 Site Instrumentation

231 At each stream reach one cross section was selected for instrumentation, to monitor water
232 level and bank activity information. Along the instrumented cross section 2-3 monitoring wells
233 were put in place. Wells at the Winooski and Lewis Creek instrumented sites were placed to
234 depths of 4.3-4.9m (14-16ft) and 3.7-4.3m (12- 14 ft) respectively; allowing monitoring water
235 levels at or above base flows. Water levels at each well were measured by use of a set of
236 Solonist level loggers, data logging pressure transducers. Data collected with these would then
237 be adjusted to account for barometric conditions, monitored by a barologger, to determine stream
238 and ground water levels. A cross section of a typical instrumented site can be seen in figure 3.2.

239 In addition to the monitoring ground water elevations it is also desirable to know the
240 stream water elevation for modeling purposes. To do this a pressure transducer was located with
241 direct exposure to the stream water. Due to the variation in the streams monitored different
242 approaches were taken to expose the logger directly to the stream. Due to depth and flow
243 conditions at the Winooski River reach directly exposing the logger to the stream water was not
244 possible. To determine the water levels in the stream without a media a non-vertical monitoring
245 well with a zenith angle of 40° used. Placements of this well, in a steep vegetated bank adjacent
246 to the stream, allowed exposure to the lower well sections directly to the stream. Shallower
247 angles of the streambanks located at the Lewis Creek reach prevented stream water monitoring

248 using a similar fashion. To expose the pressure transducer to the streams water a section of well
249 piping was capped at either end and then pinned to the streambed by use of 61cm (24in) long by
250 13cm (0.5in) diameter rebar staples. The logger was then tethered to the pins using a 0.48cm
251 (0.19in) stainless steel cable and cable crimps to prevent loss if damage was sustained to the well
252 tube.

253 If a bank failure were to occur during the course of this study it was preferable that the
254 time of the failure could be recorded in order to determine what conditions caused the bank to be
255 unstable. To do this a series of roller ball tilt switches were deployed along the instrumented
256 cross section. A four channel third party Hobo Data loggers encased in a weatherproof box was
257 then linked to the tilt switches imbedded into the streambank. A series of tilt switches were
258 placed horizontally in the bank at an off position and a 2hour logging interval (see figure 3.2.).
259 If a switch moved out of plane by more than 15° the switch would close and the logger's impulse
260 signal, 2.5 volts, would be registered, indicating a failure.

261

262 **4. Site Investigations**

263 4.1 Soil Composition

264 Typically two boreholes were augured at each cross section to a depth below the bottom
265 of the stream bed or the maximum depth possible at the site. Differences in soil characteristics
266 such as color, texture, inclusions, and odor were noted for each borehole to determine if and
267 where stratified soils were found. When significant differences were seen in the soils
268 characteristic a bulk sample was collected for grain size analysis. Shelby tube samples were also
269 collected at these and used for soil strength characterization using direct shear testing. Although

270 the sampling was done using some impact which compromised the quality of the recovered
271 samples.

272 The collected bulk samples collected were brought to the University of Vermont's geotechnical
273 facility, where standard sieve analysis (ASTM 136-96) was preformed. For some soils standard
274 sieve analysis alone was not sufficient for soil classification. On a select number of these
275 samples further testing including Atterberg limits and hydrometer analysis. Banks the Winooski
276 River reach are primarily composed of silty sandy soils with little or no cohesion and a gradual
277 coarsening with depth. Soil strengths in this reach are highly variable with no relationships
278 found between friction angle and elevation or depth from surface. An overview of soils found at
279 each reach can be seen in table 4.1. In general the Lewis Creek sediments had a coarser
280 composition than those of the Winooski reach with overlaying soils comprised primarily of
281 sandy soils with some silt. Elevational stratification is found at the Lewis Creek site with a 0.3-
282 0.45m (1-1.5ft) cobble layer found below the sandy layer and on top of a marine clay layer(see
283 figure 4.1).

284

285 4.2 Unit Weights

286 Strength and unit weight properties were determined using index properties suggested by
287 the US Naval Facilities Engineering Command Headquarters and Coduto(2001)([navfac cite](#)), for
288 each characterized soil type. Index properties found the unit weight to be 16.7kN/m^3 (106.3pcf)
289 for silty sand and 16.6kN/m^3 (103.5pcf) showing quite good agreement with field testing results
290 determined through use of the sand cone method (table 4.1.).

291

292 4.3 Soil Strengths

293 To quantify the shear strength properties of soils two separate testing procedures. Insitu
294 tests were performed by use of the Iowa Borehole Shear Test (BST) apparatus. From each
295 borehole at least two sets of shear tests were conducted where changes in soil texture or color
296 was observed. Prior to beginning a BST set a thin tube Shelby tube sample 20.3cm (8in) long
297 by 6.4cm (2.5in) in diameter was taken to smooth borehole walls increasing the accuracy of the
298 test. BSTs were conducted with a range of consolidation pressures from 15kPa (2.2psi) to
299 120kPa (17.5psi) using 15kPa (2.2psi) increments with a consolidation period after each test
300 allowing excess pore water pressure to dissipate. The values of maximum shear stress were then
301 plotted against the tests normal pressure allowing for determination of the internal friction angle
302 and the cohesion of the soil. As use of the BST is still uncommon it was necessary to compare
303 results collected with this equipment against the more commonly preformed direct shear test
304 (DST). Shelby tube samples that taken prior to running the BSTs were sectioned into 2.54cm
305 (1in) tall samples and used in a series of DSTs in accordance with ASTM D3080-04 using a Geo
306 Comp ShearTrac II automated shear device. Values obtained from both of these tests showed
307 that the friction angle values obtained using the BST were consistently higher than those
308 determined using the DST. The variation between these results was attributed to sample
309 disturbance during collection and preparation of the DST samples. As such friction angles
310 obtained by use of the BST were used for analysis purposes.

311 Strength testing of the soils found at the Winooski River site revealed highly variable with friction
312 angle values ranging from 23-43° with an average of 33° and no correlation to the elevation of
313 sampling (see figure 4.2.). Friction angles across the Lewis Creek reach were less variable with
314 angles ranging from 32-41° and an average friction angle of 39.5°(see figure 4.3.). Index

315 properties for the primary soil types at each reach had reasonable agreement with the values
316 measured. Internal shear angles of 27-37° was determined for silt and 28-38° for sand. This
317 showed good agreement with the values found at the Winooski with averaged for both indexed
318 values and measured at 33°. Friction angle values measured at the Lewis Creek reach were
319 found to be outside of range of index values.

320

321 4.4 Tensile Reinforcement

322 To quantify the effect of roots on the shear strength for soils two different approaches
323 were taken in this study. The first was to use the modified tensile reinforcement equation(3)
324 developed by Wu et al(1979). To use this equation two values needed to be determined the
325 tensile strength of the roots present at the sites, and the root area ratios in the impregnated soil.
326 Along both sites the dominant species was golden rod (*S. Canadensis*). Root balls of the plants
327 were taken at each of the study reaches and root samples less than 1 mm in diameter were
328 harvested for tensile testing. To maintain the tensile strength that would be displayed by roots in
329 the field the samples remained in a moist environment to prevent dehydration. The ends of the
330 sample root sections were anchored into epoxy molds where epoxy was poured and allowed to
331 fully harden. The samples and the epoxy anchors were then transferred to a tensile testing
332 machine where they were loaded until the root material ruptured. The average rupture force for
333 each root was then normalized by the cross sectional area of each root sample and averaged to
334 determine the tensile strength for the roots found in the samples.

335 The second method of determining the addition cohesion due to roots was by use of a
336 series of DSTs. Samples were collected from depths of up to 0.46m (1.5ft) with zenith angles of

337 0,45, and 90° denoted as β (Figure 4.4). Three samples from each angle were run as consolidated
338 drained tests with normal pressures of 13.8-48.3kPa (2-7psi). A shear stroke of 13mm (0.5in)
339 was selected to allow for full mobilization of the roots shear strength addition. The maximum
340 shear strength for each sample was then used to determine c' and ϕ' for each value of β . The
341 change in cohesion between these samples and similar samples without roots was taken to be the
342 additional cohesion due to tensile root reinforcement.

343 The samples run in the DSTs were then processed to remove the roots for calculation of
344 the root area ratio. Samples were washed through number 16 (1.18mm, 0.0464in) and 50
345 (0.30mm, 0.0118in) sieves to allow soil and roots to be separated. The roots then received a
346 second washing on the number 50 sieve where they were collected. These samples were then
347 allowed to dry at 22°C (72°F) until dry. The low drying temperature and the relatively short time
348 insured that the roots did not fully dehydrate which may change the measured root volume. Root
349 samples were observed under a magnifying glass to determine if any residual moisture remained
350 on the surface of the roots. Samples were then placed in a 50ml (1.7oz) graduated cylinder with
351 15ml (0.51oz) water, and submerged using a calibrated rod of a known diameter. The volume
352 addition of the rod was subtracted from the final additional volume in the graduated cylinder
353 allowing for the volume of the roots to be determined. Using the original volume of the sample
354 the additional cohesion due to root strength was easily calculated.

355 Cohesion due to root addition using the modified equation developed by Wu et al.(1979)
356 found to be 165-765kPa (24-111psi) using the averaged measured root tensile strength of
357 50MPa (7200psi).This unrealistic value of reinforcement calculated by this method was
358 attributed to the difficulty of measuring the root area ratio in small samples and discarded from
359 use. The values of additional cohesion determined by use of the DSTs was found to be 6.3kPa

360 (0.92psi) with no noticeable dependence on the zenith angle of sampling (see figures 4.5. and
361 4.6.).

362

363 4.5 Matric Suction

364 A series of tensiometer measurements were conducted during the summers of 2007 and
365 2008. Tensiometer were placed at various depths, 6-3 in (0.15-0.81m), below the ground
366 surface and allowed to equilibrate for 5hrs. Throughout the course of this sampling only one
367 negative pore-water pressure value was measured. This indicates one of two things either the
368 moisture level of the ground was high enough to have minimal suction values or proper
369 installation of the tensiometers was not conducted.

370

371 4.6 Erosion Characteristics (test not yet conducted)

372 Characterization of the soils erosion properties “were” carried out by the use of a series
373 of Jet Erosion Test, in accordance with ASTM standard D 5852-00. Samples measuring 0.25m
374 (10in) in diameter by 0.20m (8in) were collected using a sharp edged polyvinyl sampling tube
375 from the base material at each stream reach and transported to the University of Vermont’s
376 geotechnical testing facility. Prior to running the test the soil sample was extruded and leveled to
377 provide an initially uniform surface. Scour depth was measured at 10 min increments for the
378 duration of the test. Each test was run until the equilibrium scour depth was reached. The soil
379 was then advanced upward in the tube and retested. Results from the test were then used to
380 determine the coefficient of erosion and the critical shear velocity. Average erosion rates were

381 determined to be ### and ### for the Winooski River and Lewis Creek reaches respectively.
382 While the critical shear stresses were found to be ### and ### for the Winooski River and Lewis
383 Creek reaches respectively.

384

385 **5. Analysis**

386 5.1 Lateral Retreat

387 The point of maximum lateral retreat was noted for each cross section along both of the
388 reaches and then averaged. Across the seven sites at the Winooski average maximum lateral
389 retreat rate was measured to be 0.061m/month (0.2ft/month). The average rate at the Lewis
390 Creek site was found to be similar at 0.046m/month (0.15ft/month). Although the mechanism
391 that caused these rates is unknown removal on steeper banks were from areas not commonly
392 exposed to water indicating a mass failure due to instability. Shallower banks with material
393 removal were generally from the toe of the slope and in a vertical fashion indicating scour
394 erosion as a cause. Additionally two cross sections of the Lewis Creek reach, located at the start
395 of a cut bank showed signs of aggregation towards the base of the slop with no indication of a
396 failure that would have deposited these materials.

397

398 5.2 Slope Stability

399 Several models were used to look to evaluate the stability of the streambank cross
400 sections, the widely available toe erosion and slope stability spread sheet package, GeoStudio's
401 stability modeling program SLOPE/W, and efforts were made on merging several algorithms
402 into a single model. Initial modeling efforts were focused on the toe erosion and slope stability

403 spread sheet package which is capable of modifying the bank geometry of a slope if the erosion
404 characteristic and the size of the hydraulic event is known. Unfortunately the coupled slope
405 stability model in this could not generate curved slip surfaces similar to those seen in the field.
406 The most versatile of these modeling programs was the GeoStudio's SLOPE/W program. In this
407 modeling package the phreatic surface determined from the pressure transducers could easily be
408 entered. SLOPE/W was capable of creating and analyzing multiple slip surfaces within a user's
409 specified range of possible slip surfaces. Due to the ability of this program to generate multiple
410 failure surfaces it was used for the bulk of stability analysis performed. Using information
411 gathered from the pressure transducers 10 points in time were selected for analysis at each cross
412 section.

413 As in the case of the Winooski River site a full cross sectional survey was not possible
414 which would allow for easy computation of the changes in stream elevation across all of the
415 sites. To compensate for this difficulty several water elevation surveys were conducted during
416 base flow and one flow event exceeding 80% percentile flow for the stream. Linear interpolation
417 and extrapolation was then used to offset the stream to the appropriate elevation at a given cross
418 section to be modeled.

419 To date slope stability analysis has been performed on three sets of data collected for
420 analysis of 90 streambank situations. Streambank geometries of the Lewis Creek Reach were
421 modeled and then analyzed for stability using both the soils measured and index properties. Due
422 to the highly variable strength characteristic found at the Winooski River reach slope stability
423 analysis has only been performed using the determined index properties. Further analysis will be
424 performed on the Winooski cross sections using section specific measured properties. Table5.1.

425 shows an overview of the results for the differences in satiability for the Lewis Creek site using
426 index and measured soil properties

427 Using measured soil properties two sites in the study were found to have a factor of
428 safety over one during at least one time step. With highest factor of safety observed for any
429 given sight was 1.04 at the instrumented site on the Lewis Creek reach, indicating that this was
430 the most stable site in the study. The remainder of the sites was found to have factors of safety
431 to be less than one for all of the time steps analyzed. Using indexed properties found factors of
432 safety consistently lower by approximately 20%. The factors of safety relating to water
433 elevation for each analysis scheme can be seen in the following figures 5.1., 5.2., and 5.3.

434

435 **6. Conclusions**

436 1- Results from the BSTs and DST s showed that results obtained were comparable with each
437 other for silty sands and sandy silts. Results from the BST obtained slightly higher values of the
438 soils internal friction angles than those found by the DST which may affect the determination of
439 a slopes stability.

440 2- The factors of safety observed at 11 of the 13 sites that were analyzed using classical soil
441 mechanics were found to be less than one for all of the water levels that were observed, showing
442 that using classical soil mechanics is inappropriate for the stream reaches and soil types
443 analyzed.

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447 **7. References**

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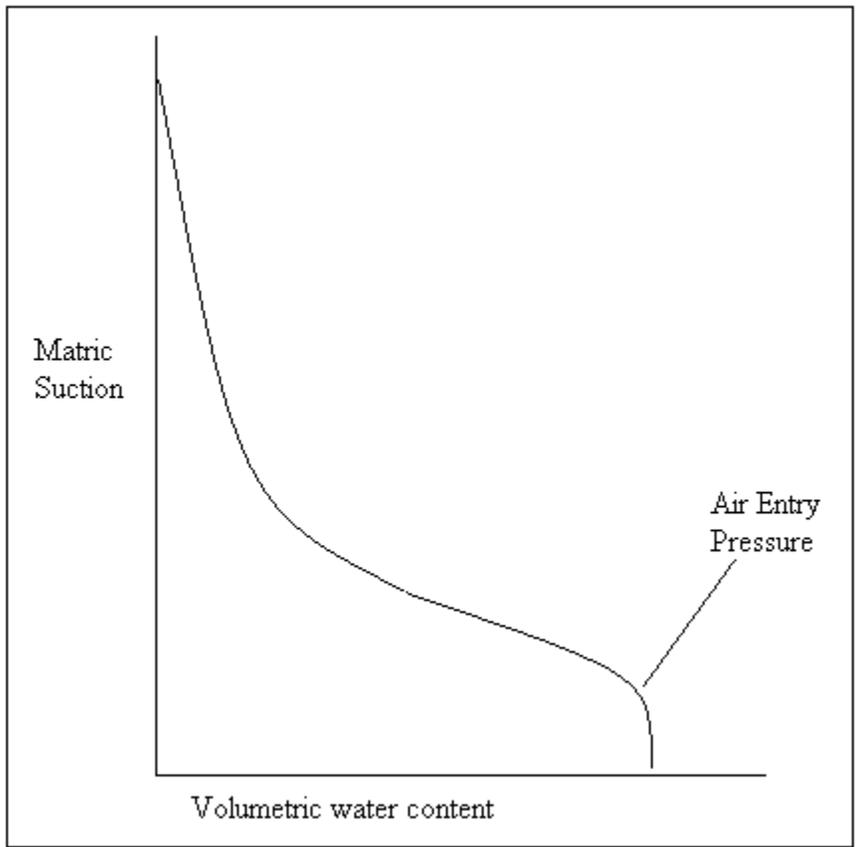
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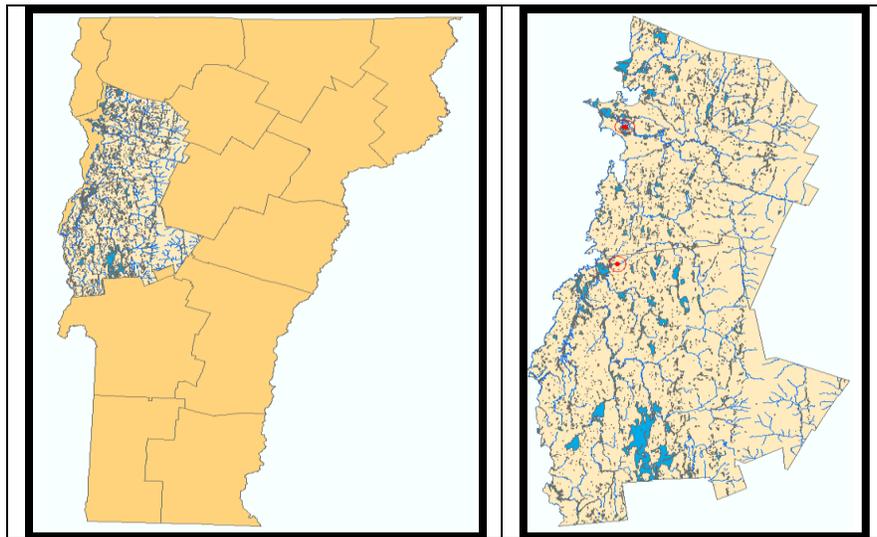
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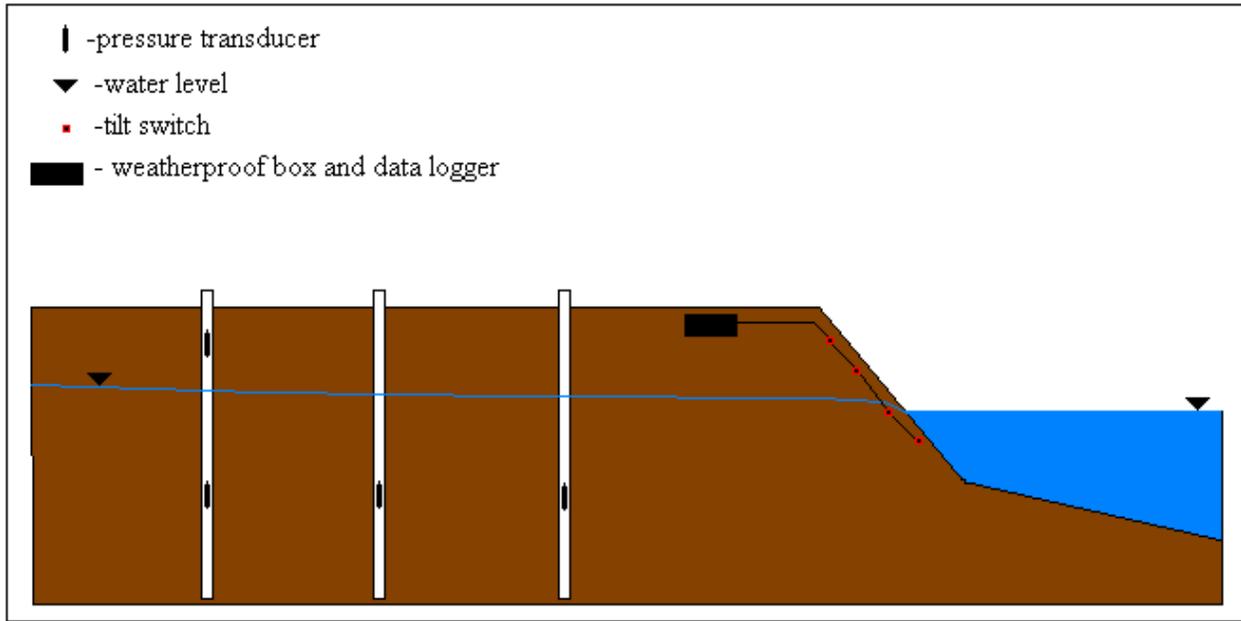
528 Fig.2.1.



529 Fig.3.1.

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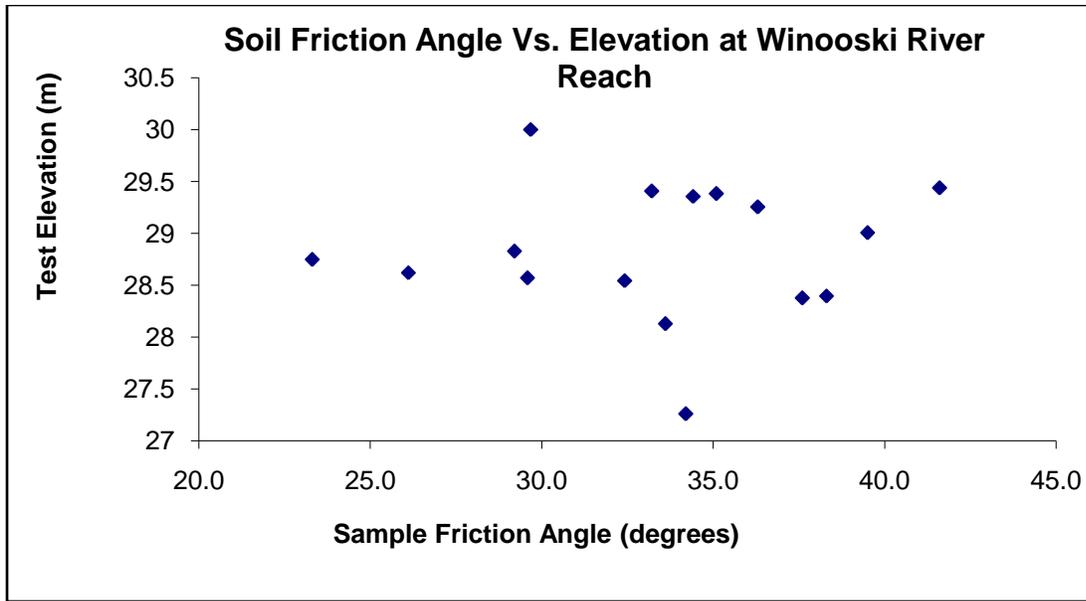
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533 Fig. 3.2



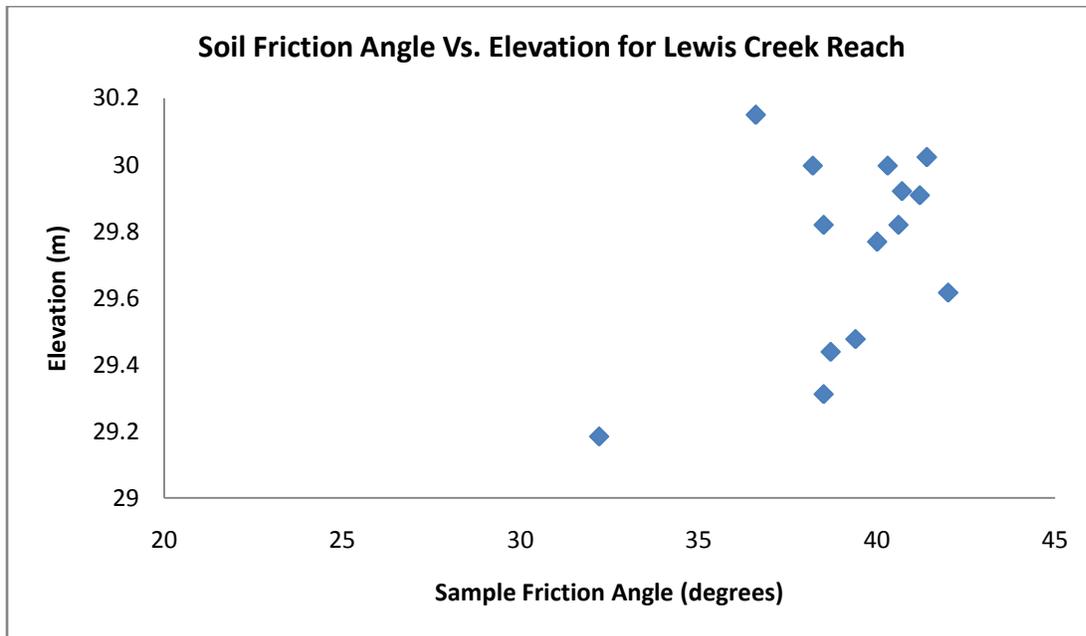
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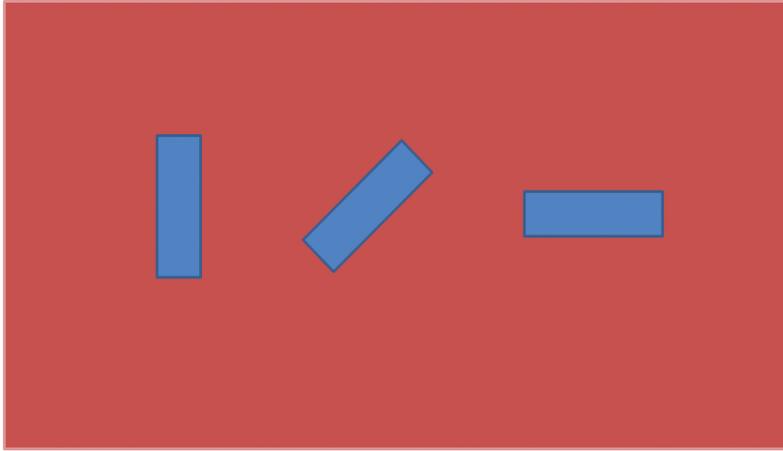
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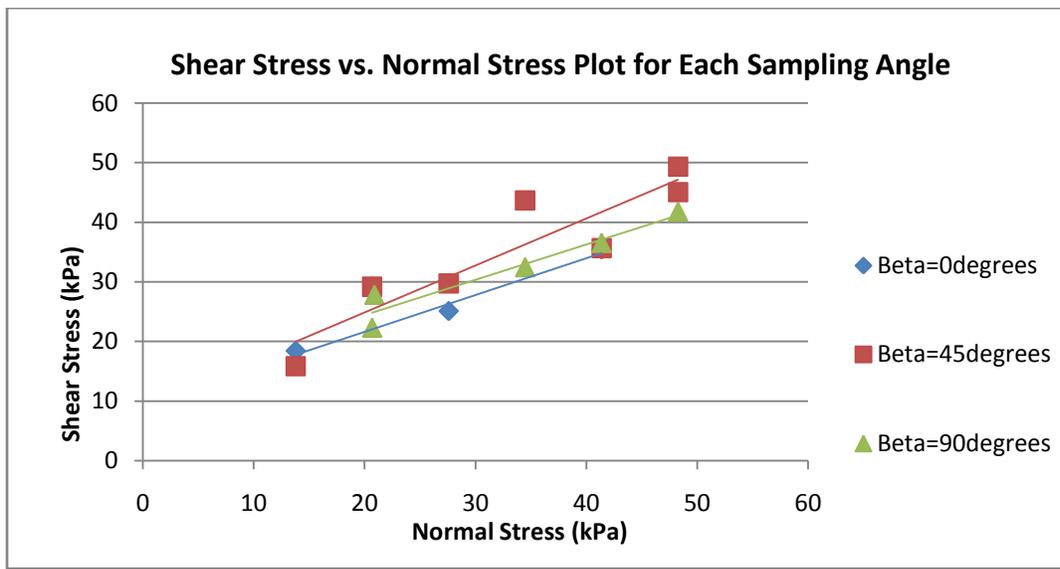
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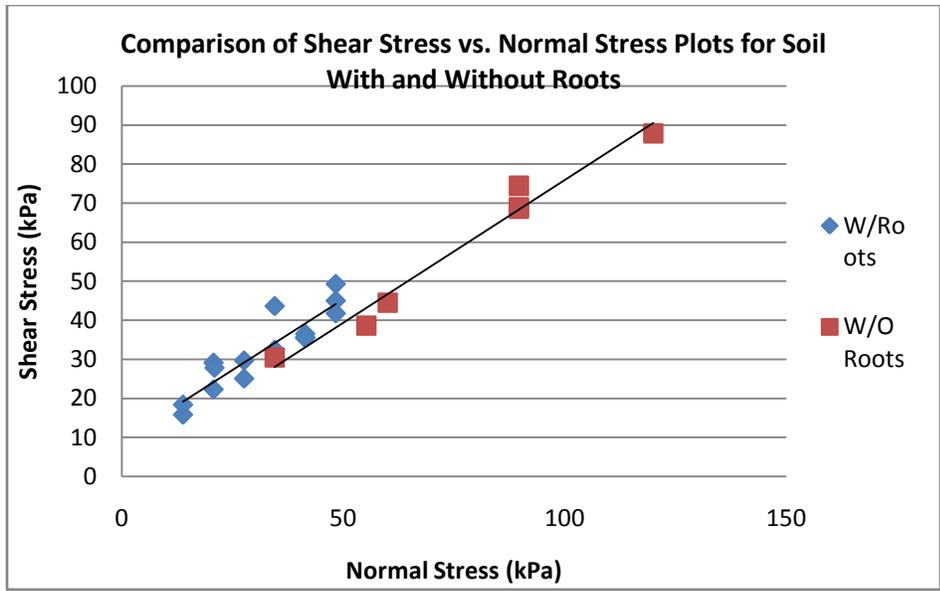
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541 Fig. 4.4.



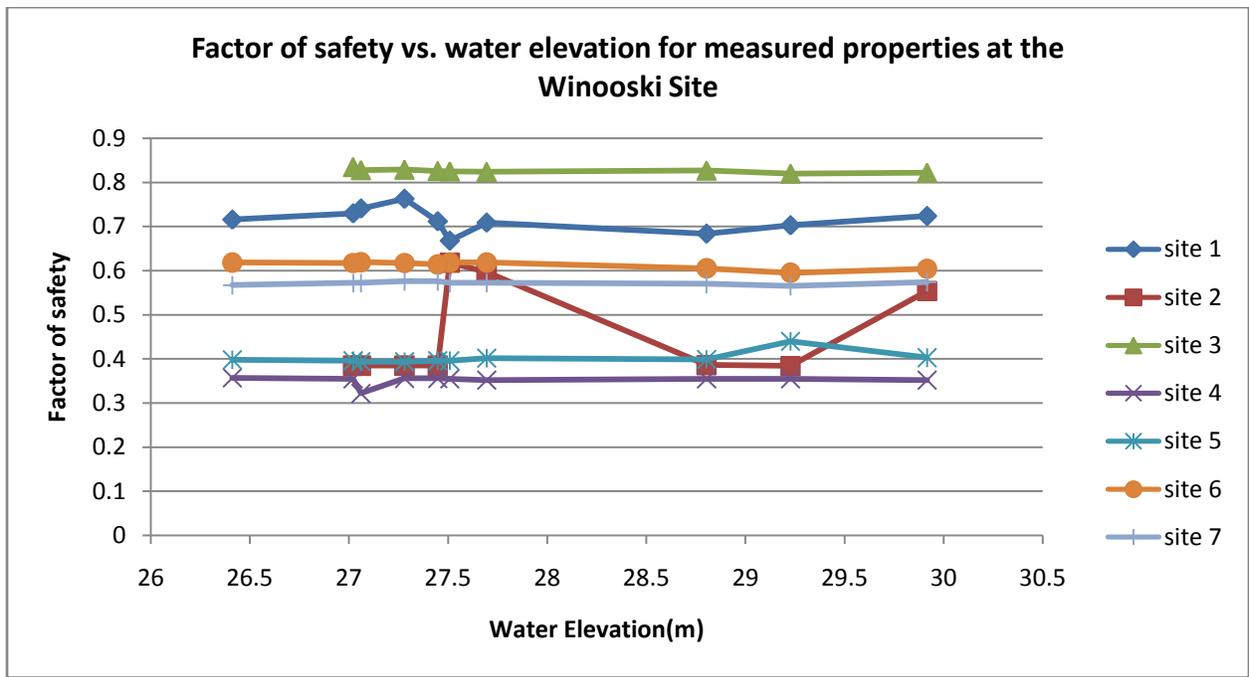
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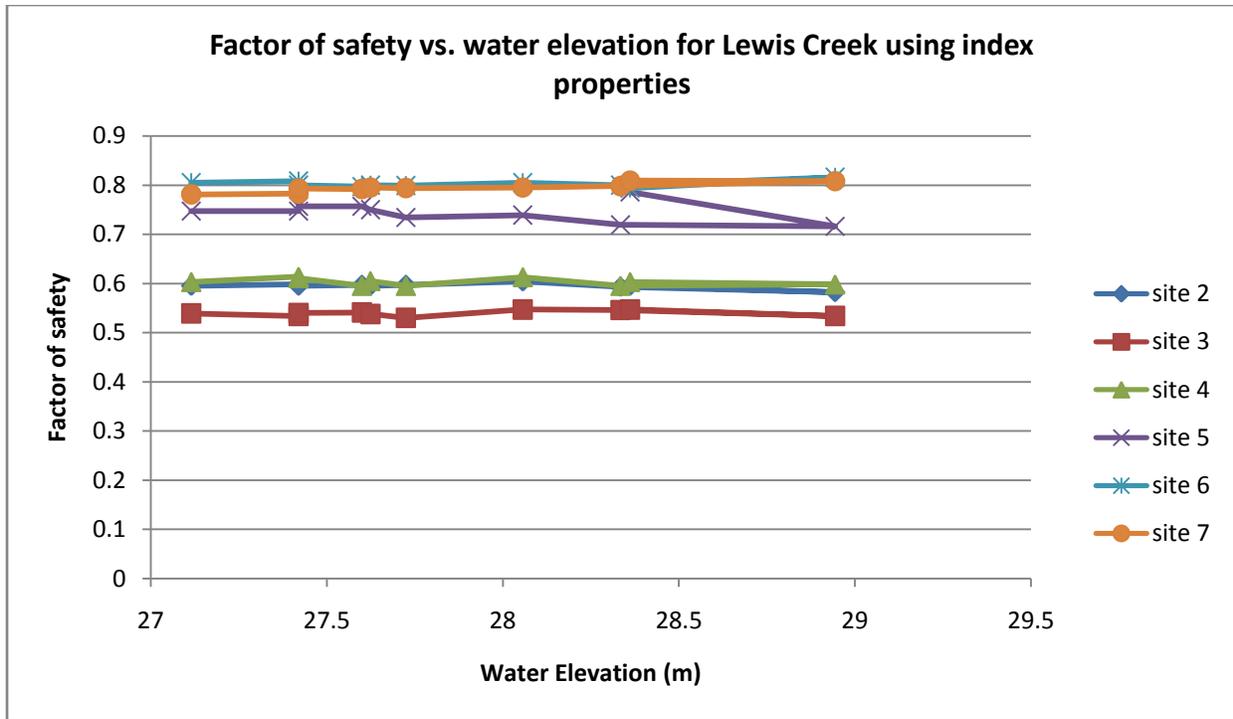
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545 Fig 4.6.



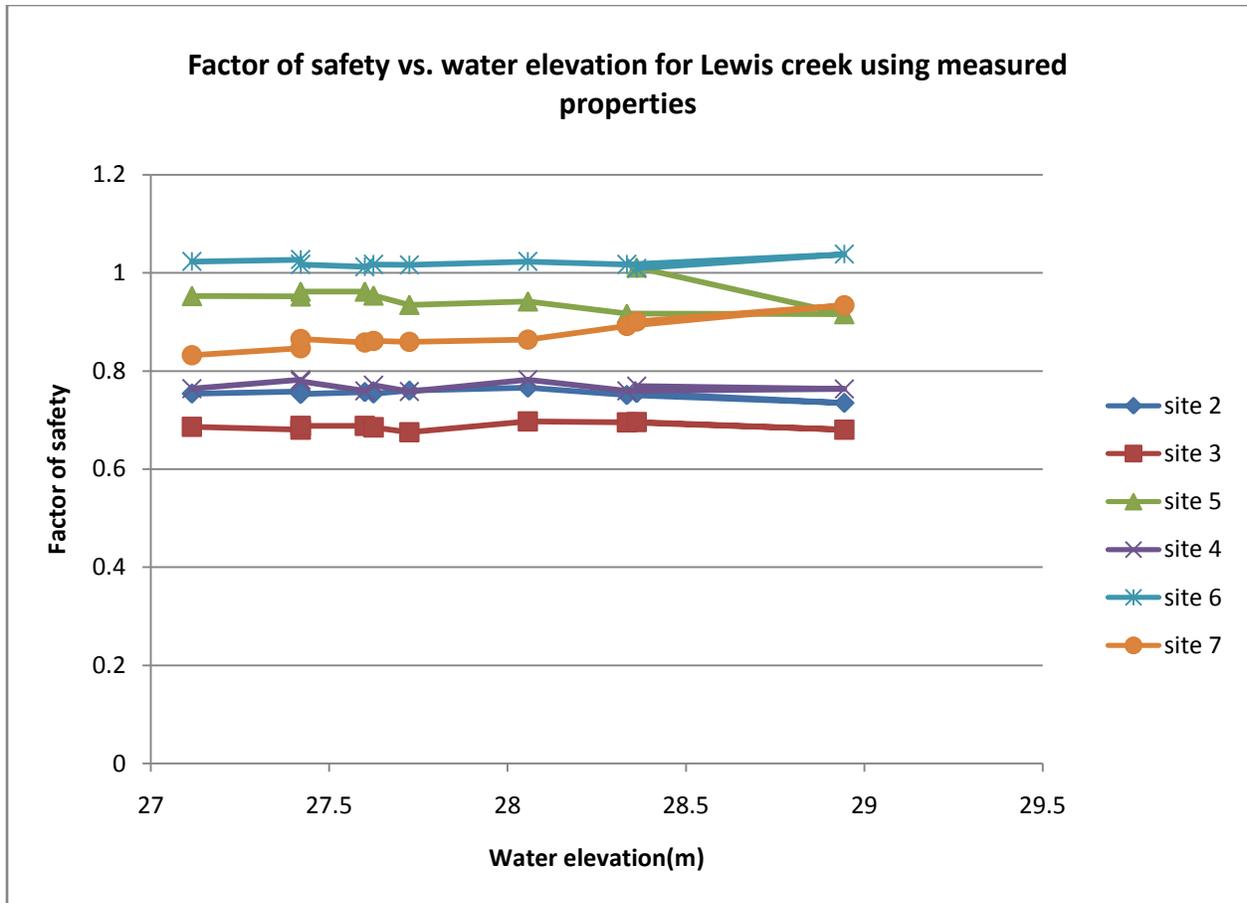
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547 Fig. 5.1.



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549 Fig. 5.2.



550

551 Fig 5.3.

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553 Fig.2.1. Conceptual drawing of the soil water characteristic curve. Note air entry pressure is the
 554 transitional point between saturated and unsaturated conditions.

555 Fig.3.1. Location of the two studied reaches

556 Fig.3.2 Site instrumentation scheme used to collect water level and bank activity data

557 Fig.4.1. Typical stream bank selected for observation. This image of the Lewis Creek
 558 instrumented cross section displays stratification common to this reach.

559 Fig.4.2. Plot of soil friction angle across elevations encountered at the Winooski River reach.

560 Fig.4.3. Plot of soil friction angle across elevations encountered at the Lewis Creek Reach

561 Fig.4.4. Depiction of sample angles for root shear testing. Left to right angles of $\beta= 0, 45,$ and
 562 90 respectively.

563 Fig.4.5. Shear strength testing results for root impregnated samples collected at various angles.
564 Tight clustering shows the limited affect of sample angle on shear strength.

565 Fig.4.6. Comparisons of soils shear strengths with roots present to those without. The vertical
566 offset is representative of the additional cohesion provided by roots.

567 Fig.5.1. Determined factors of safety for the Winooski reach at each water elevation analyzed

568 Fig.5.2. Determined factors of safety for each of Lewis Creek sites at all 10 timesteps using
569 indexed.

570 Fig.5.3. Determined factors of safety for each of the Lewis Creek sites for each time step using
571 measured soil properties at each water elevation analyzed

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592 Table 4.1. Site information and soil properties for both the Winooski River and Lewis Creek
593 sites. Note Friction angles for Lewis Creek Site are only SM soils.

Site	Properties				
	USCS soil classes	Dry unit weight (kN/m ³)	Water Content	Friction Angle (degrees)	
Range				Average	
Winooski River	SM,ML,MH	17.1	0.3	22-42	33
Lewis Creek	SM,CH,	17.2	0.22	36-43	39.5

594

595

596 Table 5.1. Average factors of safety found at the Lewis Creek Site and percent difference
597 between measured and index properties

	Bank ID					
	2	3	4	5	6	7
Measured Properties	0.75	0.69	0.77	0.95	1.02	0.87
Index Properties	0.6	0.54	0.6	0.75	0.8	0.79
Percent Difference	21%	22%	22%	22%	21%	9%

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