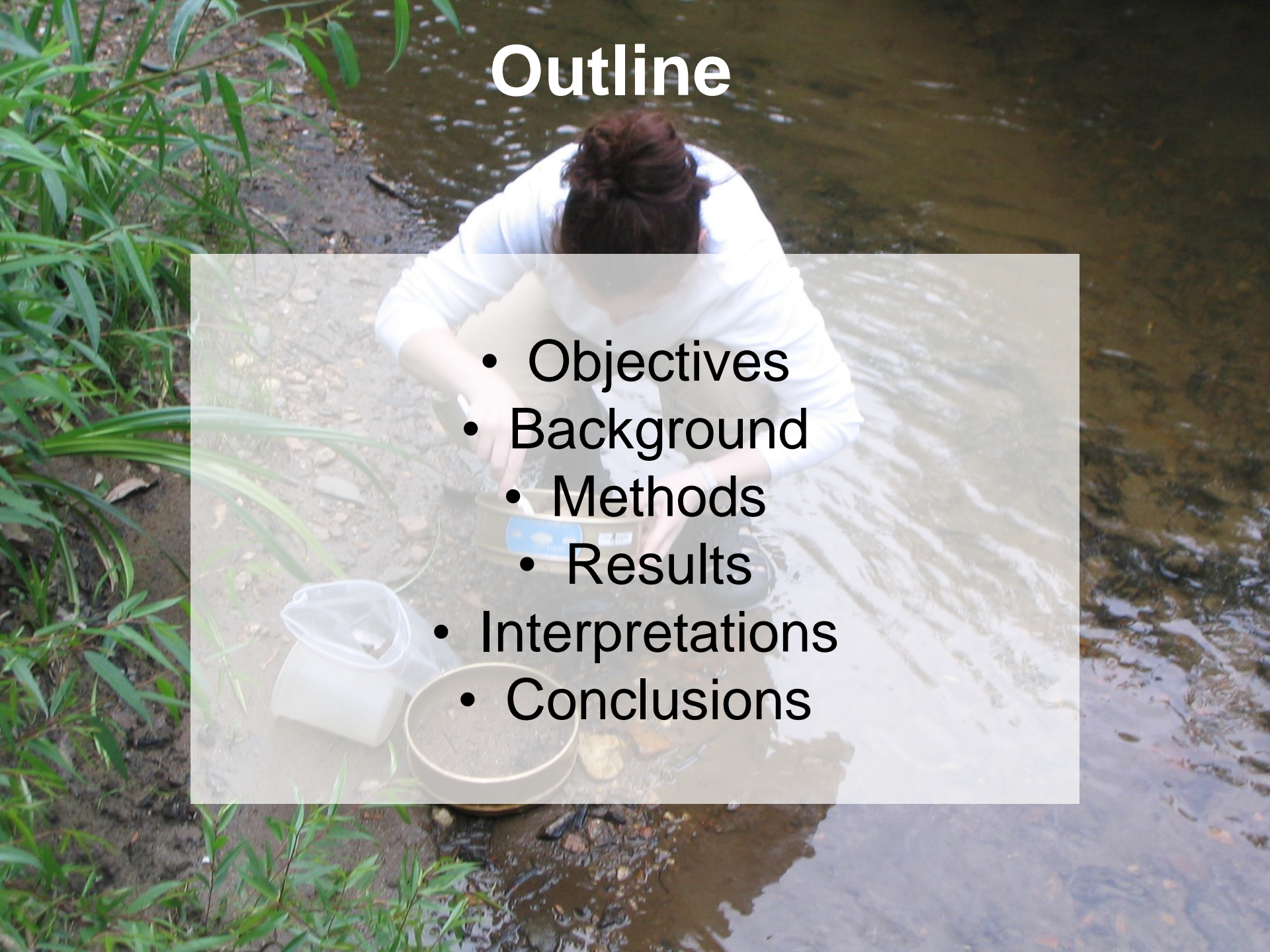


**$^{10}\text{Be}$  EROSION RATES AND LANDSCAPE  
EVOLUTION OF THE BLUE RIDGE  
ESCARPMENT, SOUTHERN APPALACHIAN  
MOUNTAINS**

**Colleen L. Sullivan  
MS Thesis Defense  
April 6, 2007  
Advisor: Paul Bierman**



# Outline

- Objectives
  - Background
  - Methods
  - Results
  - Interpretations
  - Conclusions
- 
- A person wearing a white lab coat is crouching on the bank of a stream, performing a water filtration experiment. They are using a sieve to filter water into a bucket. A white mesh bag and a white bucket are on the ground nearby. The stream is shallow and clear, with rocks visible at the bottom. The background shows green foliage and a clear sky.

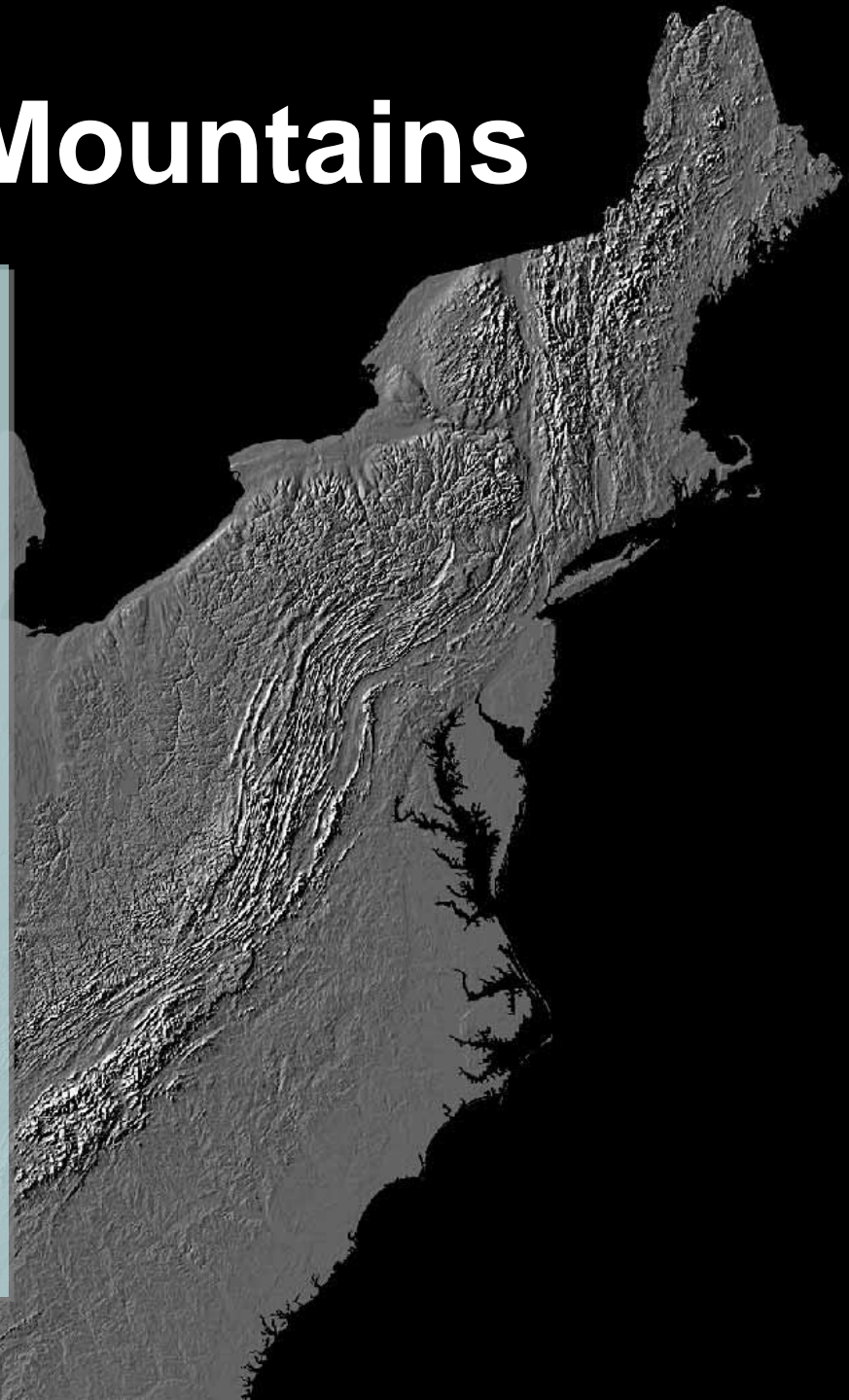


# Objectives

- to determine whether grain size influences  $^{10}\text{Be}$  concentration in fluvial sediment on and near the Blue Ridge escarpment (BRE);
- to quantify basin-scale  $^{10}\text{Be}$  erosion rates for the BRE and the surrounding provinces;
- to test for relationships between  $^{10}\text{Be}$  erosion rates and specific landscape characteristics;
- to determine whether the BRE has evolved according to a model of ongoing & parallel retreat or by a model of rapid erosion following rifting and subsequent landscape stability.

# Appalachian Mountains

- Paleozoic orogenic events created rugged mountain range
- Erosion during the Permian and Triassic
- Continental rifting and rift margin uplift in the Mesozoic (origin of BRE from rift fault)
- Followed by denudation and isostatic compensation





# The southern Appalachian Mountains

A topographic map of the southern Appalachian Mountains region, showing elevation and geological features. The map is color-coded by elevation, with green representing lower elevations and brown/purple representing higher elevations. A thick black line runs along the eastern edge of the mountains, marking the escarpment. A yellow line follows the western edge of the mountains, marking the Blue Ridge. A red line follows the southern edge of the mountains, marking the Piedmont. A large, light blue area in the southwest represents a Mesozoic Rift Basin. The map also shows major rivers and the Atlantic Ocean to the east.

Mesozoic Rift Basins

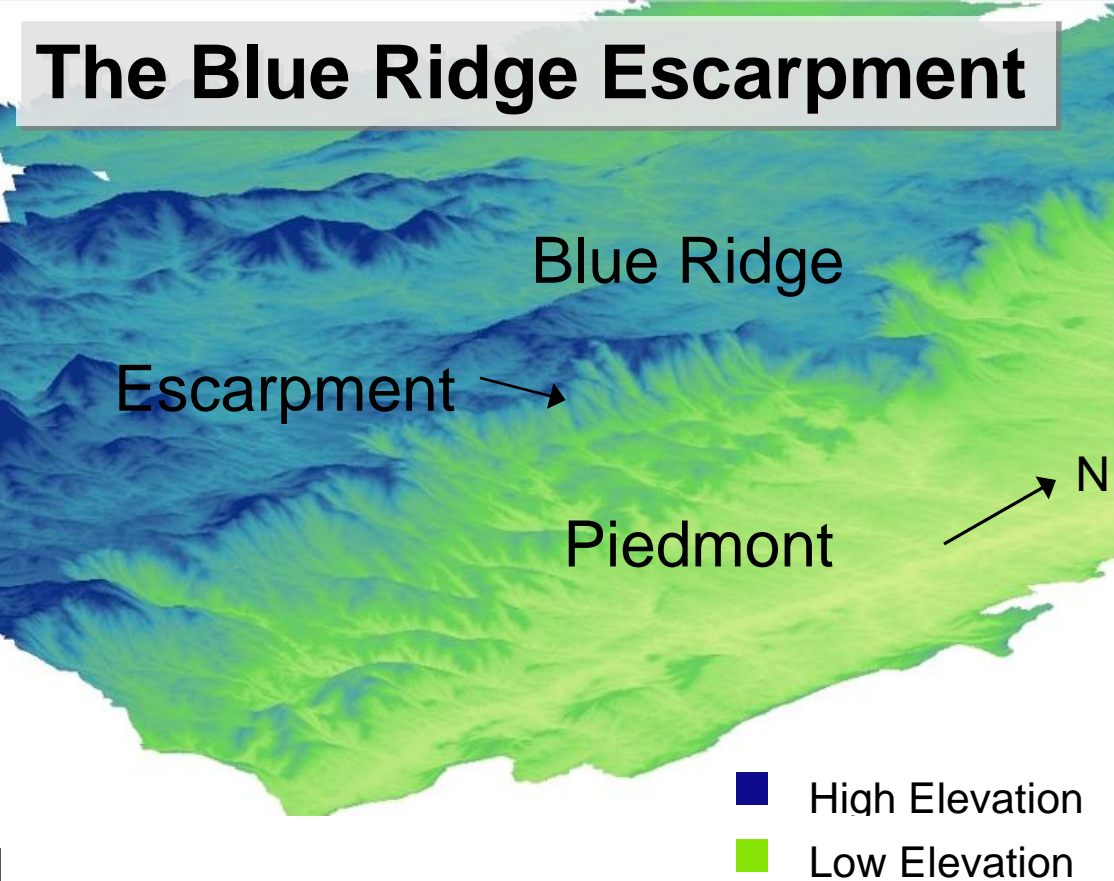
Escarpment

Blue Ridge

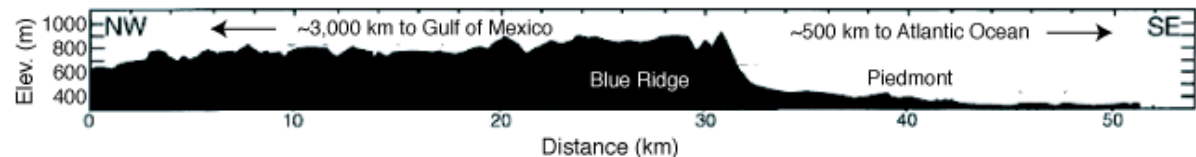
Piedmont



# The Blue Ridge Escarpment

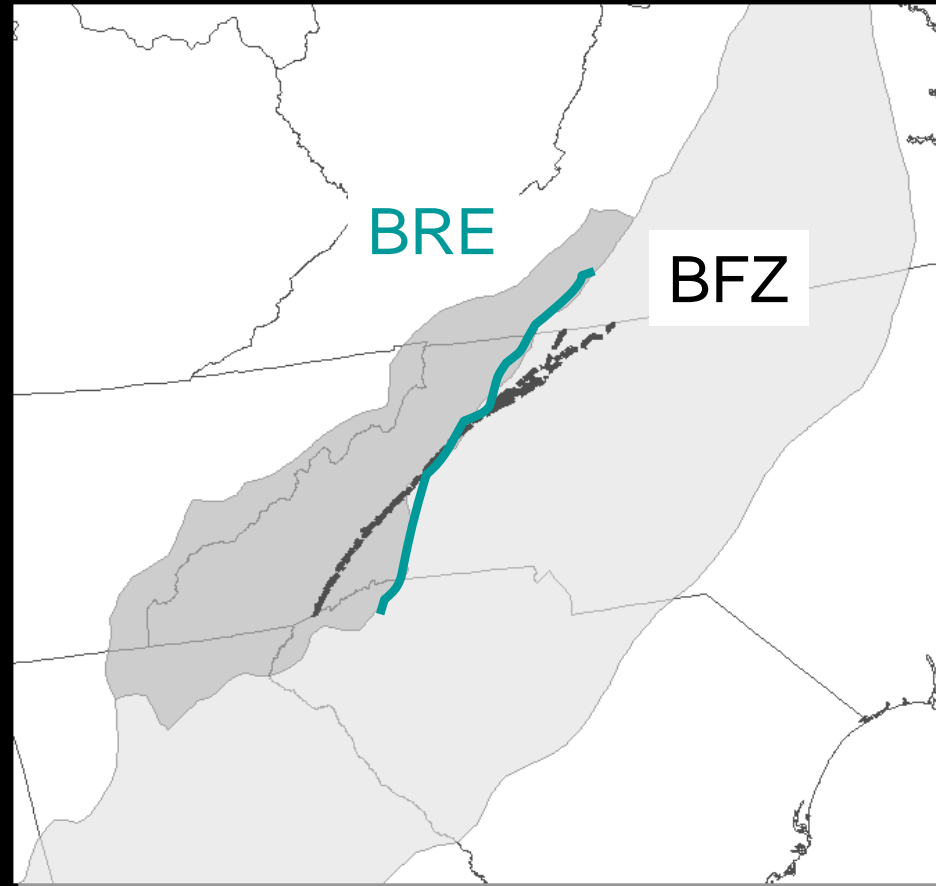


- Distinct boundary between less rugged Blue Ridge and Piedmont
- >450 km long
- Can be >500 m high
- Slopes up to 20-30°
- Asymmetric drainage divide
- Generally within lithology of micaceous schist and gneiss; thus morphology cannot be attributed to differences in bedrock erodability



# Brevard fault zone

- Oriented SW-NE
- Extends ~600 km from AL to VA
- Activated during all Appalachian orogenies, well before the rifting events that formed the BRE
- The BFZ only coincides with the BRE for 50-60 km



# Great escarpments

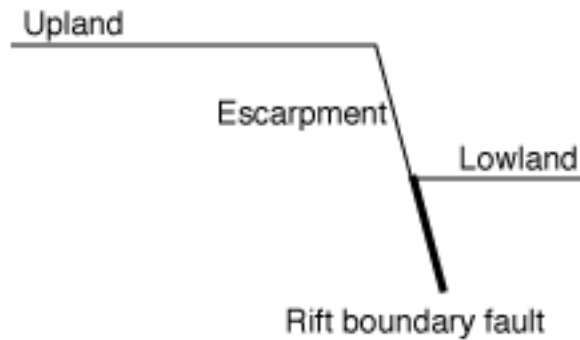


The Earth today.

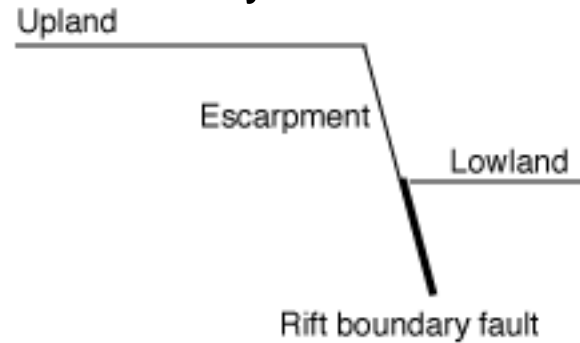


# Escarpment evolution

Ongoing & steady  
A) retreat



Significant retreat  
following rifting, then  
stability  
B)



Rifting

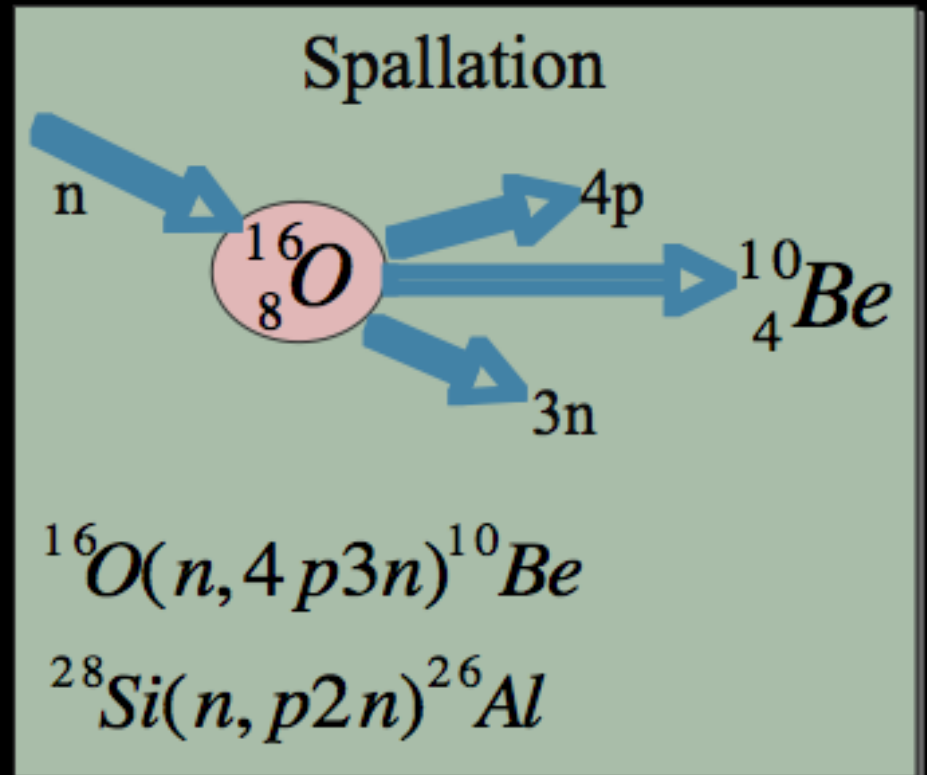


Millions of Years



# What is $^{10}\text{Be}$ ?

- $^{10}\text{Be}$  accumulates within rock that becomes sediment as it approaches surface
- Time scale of  $10^4$ - $10^5$  years





# Inferring erosion rates with $^{10}\text{Be}$

- Rivers mix sediment moving out of drainage basins, thus the concentration of  $^{10}\text{Be}$  in fluvial sediment indicates sediment production rates on basin hillslopes.
- Cosmic ray dosing as bedrock is exposed can be used to model bedrock lowering rates.

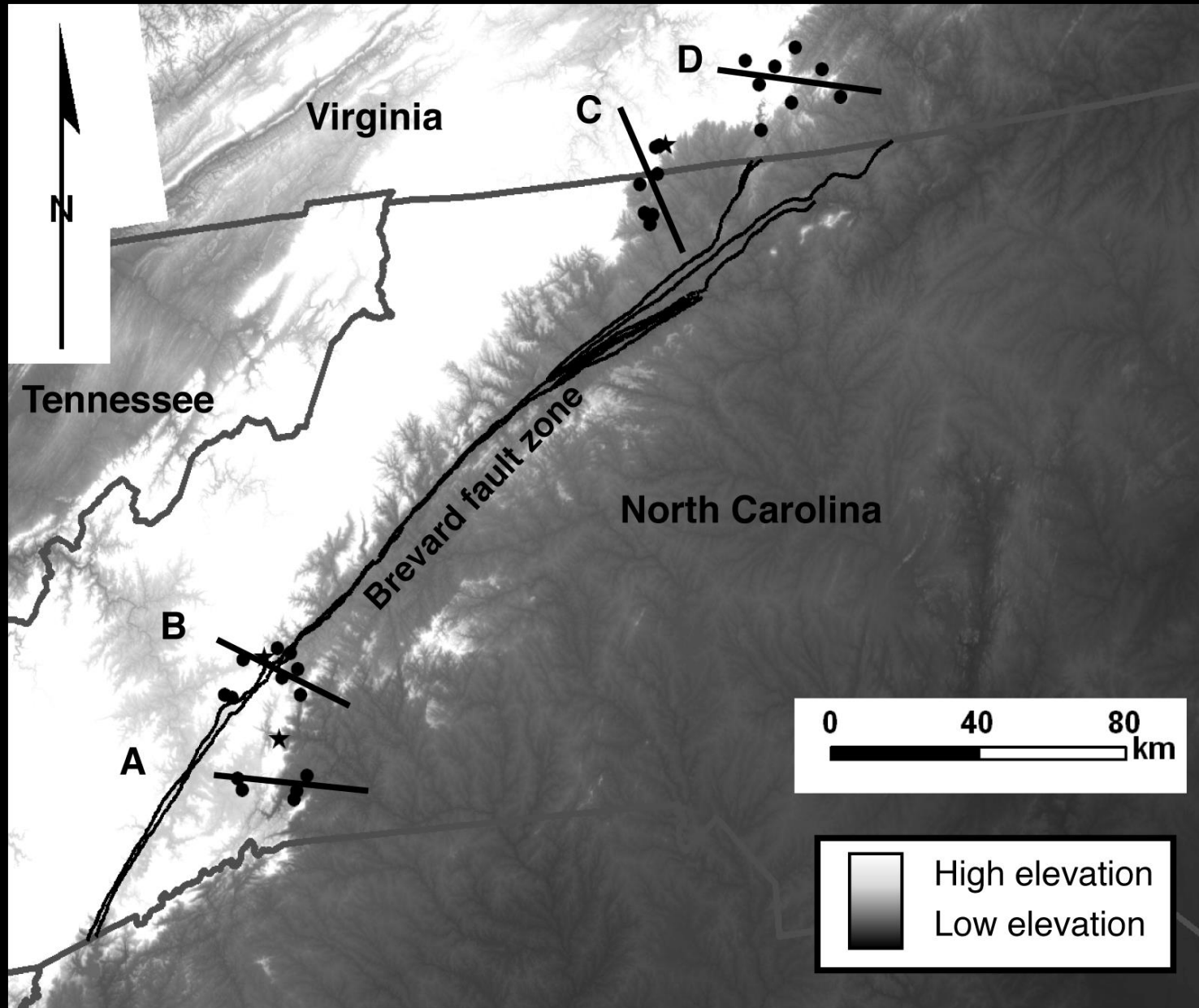


# Assumptions

- Well mixed sediment
- No inheritance from prior period of near-surface irradiation
- Sediment transport and production are in steady state



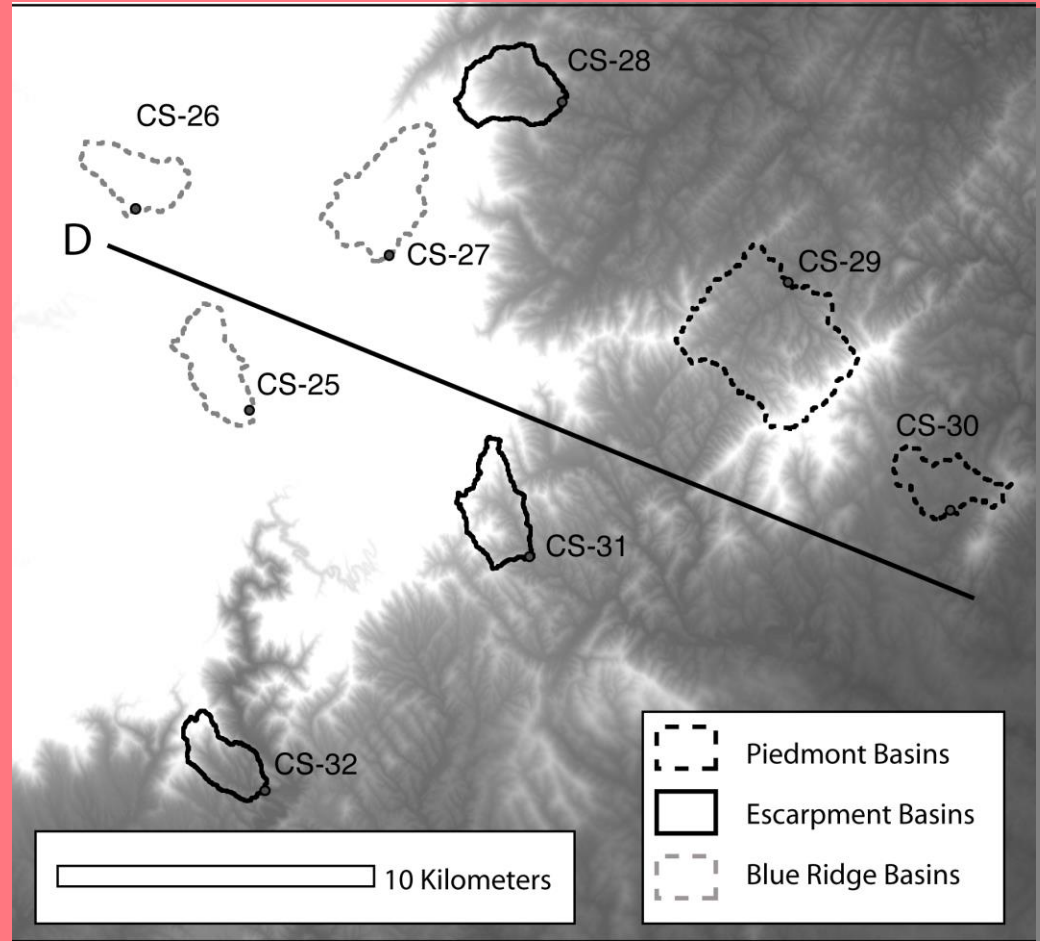
# Transect Locations





# Basin Selection

- Selected basins based on:
  - size
  - slope
  - province
- I used a GIS database to select basins for Transects B & D
- Manually selected basins from topo maps for Transects A & C



# Field methods

- Collected fluvial sediment from 32 basins:
  - Transects B & D field sieved (0.25-0.85 mm)
  - Transects A & C collected mixed grain sizes and sieved in the lab
- Collected 3 bedrock samples from escarpment





# Lab methods



- Purified quartz for 53 samples:
  - 32 basins
  - 6 grain size splits
  - 3 bedrock
- Jennifer Larsen isolated  $^{10}\text{Be}$  from all samples

# AMS at Lawrence Livermore National Laboratory

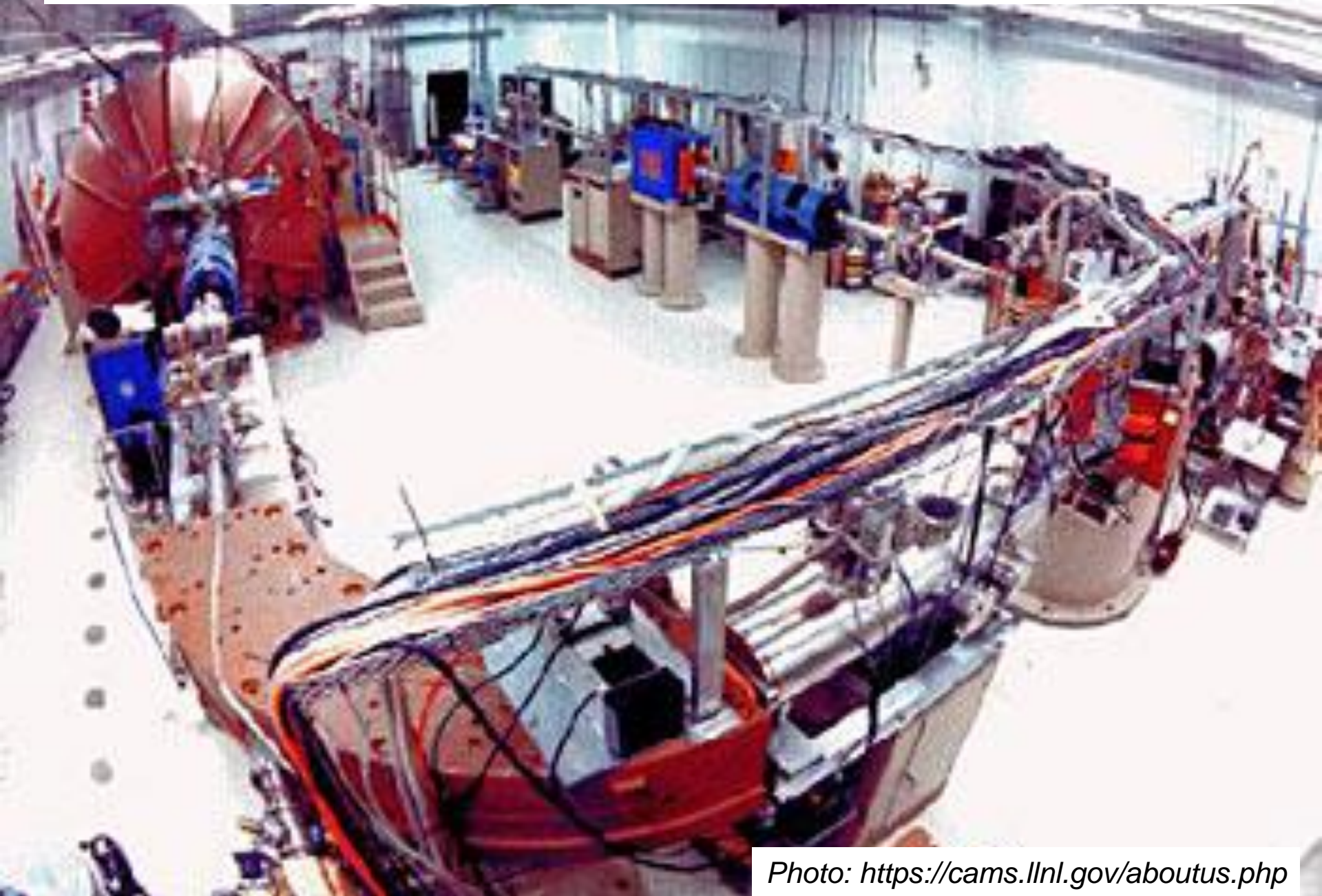
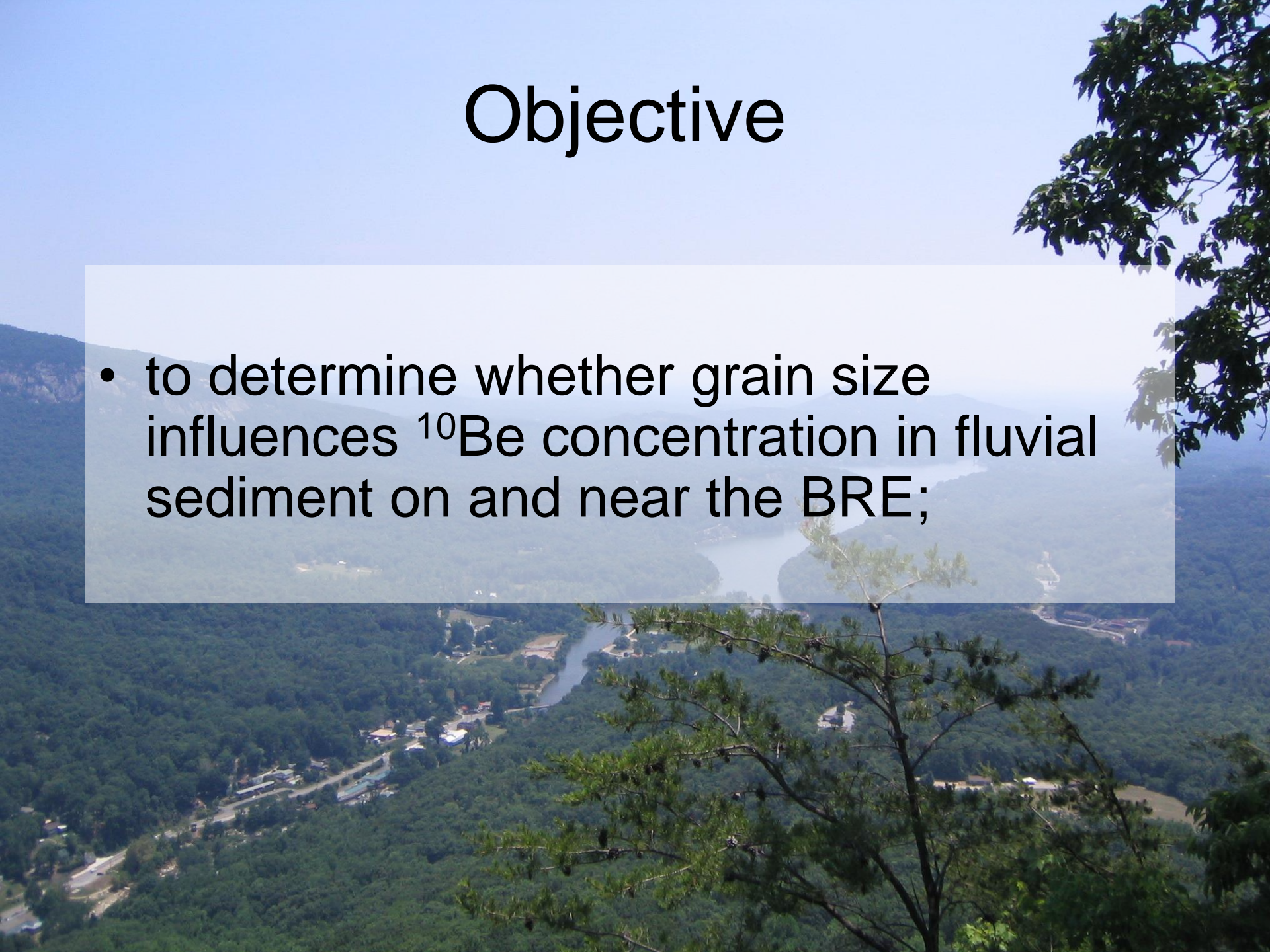


Photo: <https://cams.llnl.gov/aboutus.php>



# Objective

- to determine whether grain size influences  $^{10}\text{Be}$  concentration in fluvial sediment on and near the BRE;

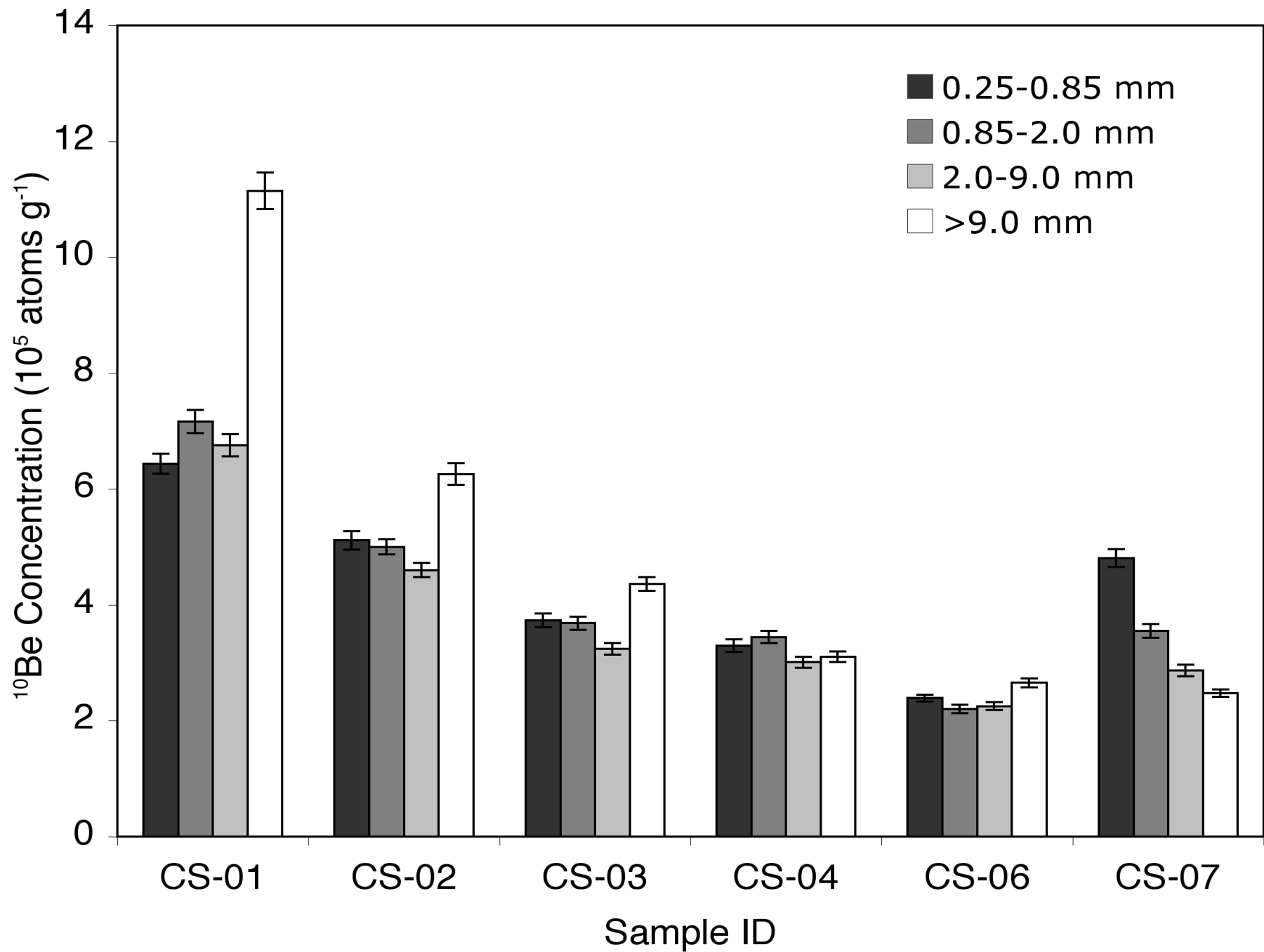




# Grain Size... Who Cares?

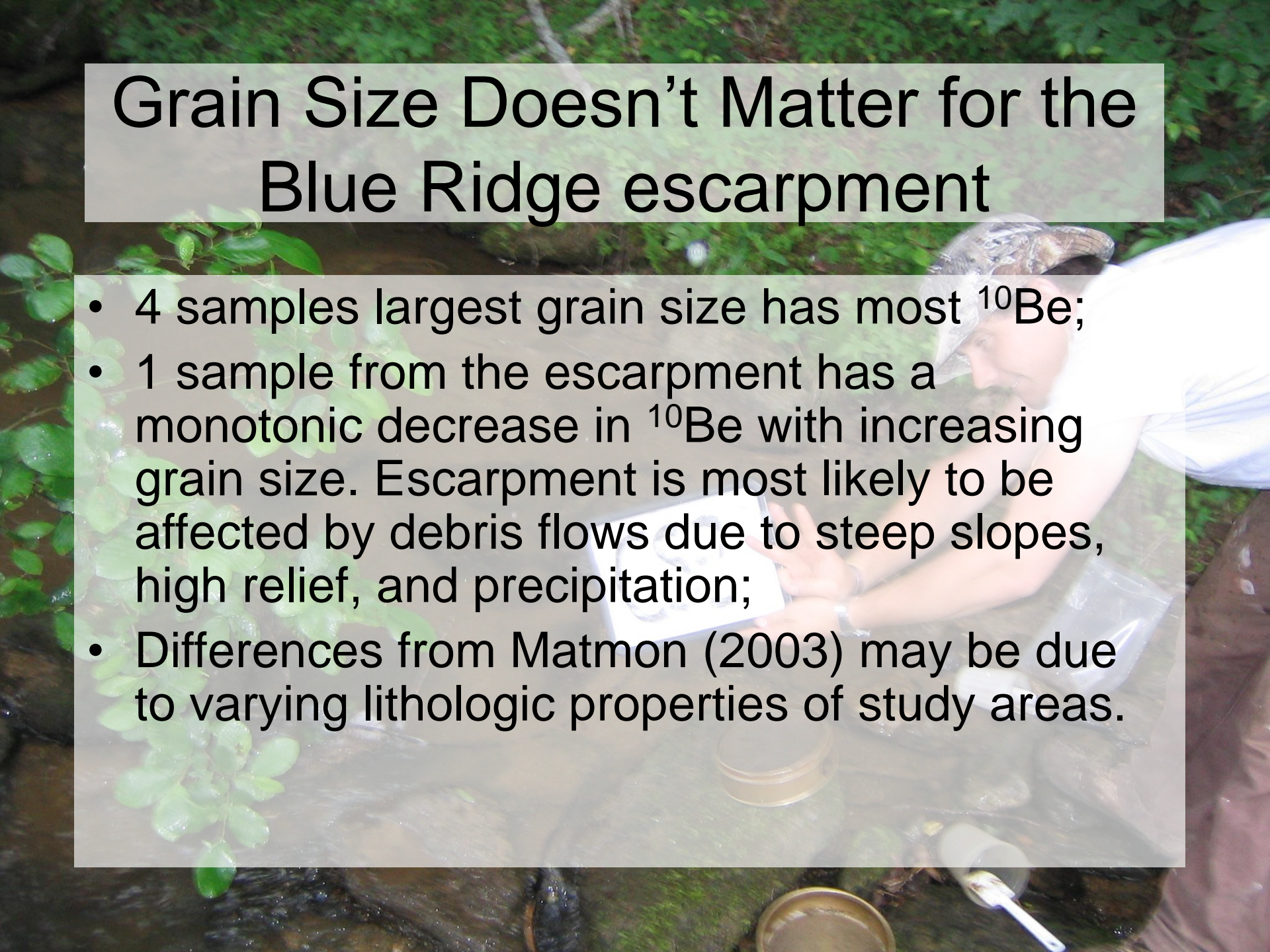
- Brown et al. (1995) suggested that lower  $^{10}\text{Be}$  concentrations in larger grain sizes could result from mass wasting events that excavate and carry previously shielded coarse material rapidly down slope.
- Matmon et al. (2003) suggested that the systematic difference in  $^{10}\text{Be}$  concentrations between small and large grains in the Great Smoky Mountains results from source area elevation and clast transport distance.





# Grain Size Doesn't Matter for the Blue Ridge escarpment

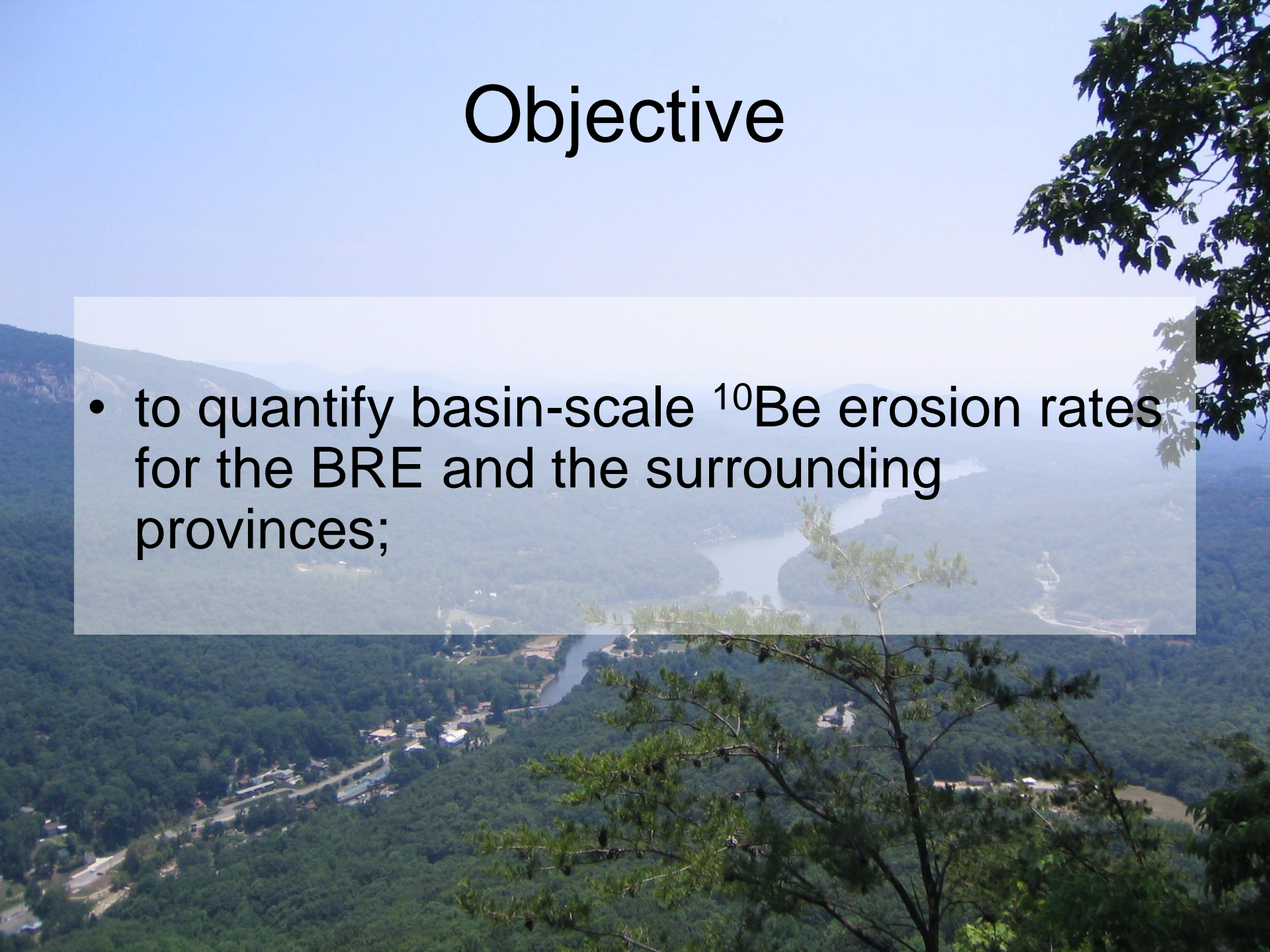
- 4 samples largest grain size has most  $^{10}\text{Be}$ ;
- 1 sample from the escarpment has a monotonic decrease in  $^{10}\text{Be}$  with increasing grain size. Escarpment is most likely to be affected by debris flows due to steep slopes, high relief, and precipitation;
- Differences from Matmon (2003) may be due to varying lithologic properties of study areas.





# Objective

- to quantify basin-scale  $^{10}\text{Be}$  erosion rates for the BRE and the surrounding provinces;





# Bedrock samples

- Heavy vegetation in June
- 3 sample sites along escarpment
- Highly variable results
  - CSB-1 (gneiss)  $56.8 \text{ m My}^{-1}$
  - CSB-2 (gneiss)  $1.7 \text{ m My}^{-1}$
  - CSB-3 (mica schist)  $17.4 \text{ m My}^{-1}$
- Lack of natural amalgamation



# Basin-Scale Erosion Rates



Escarpment (n=7)  
 $20.1 \pm 6.6 \text{ m My}^{-1}$

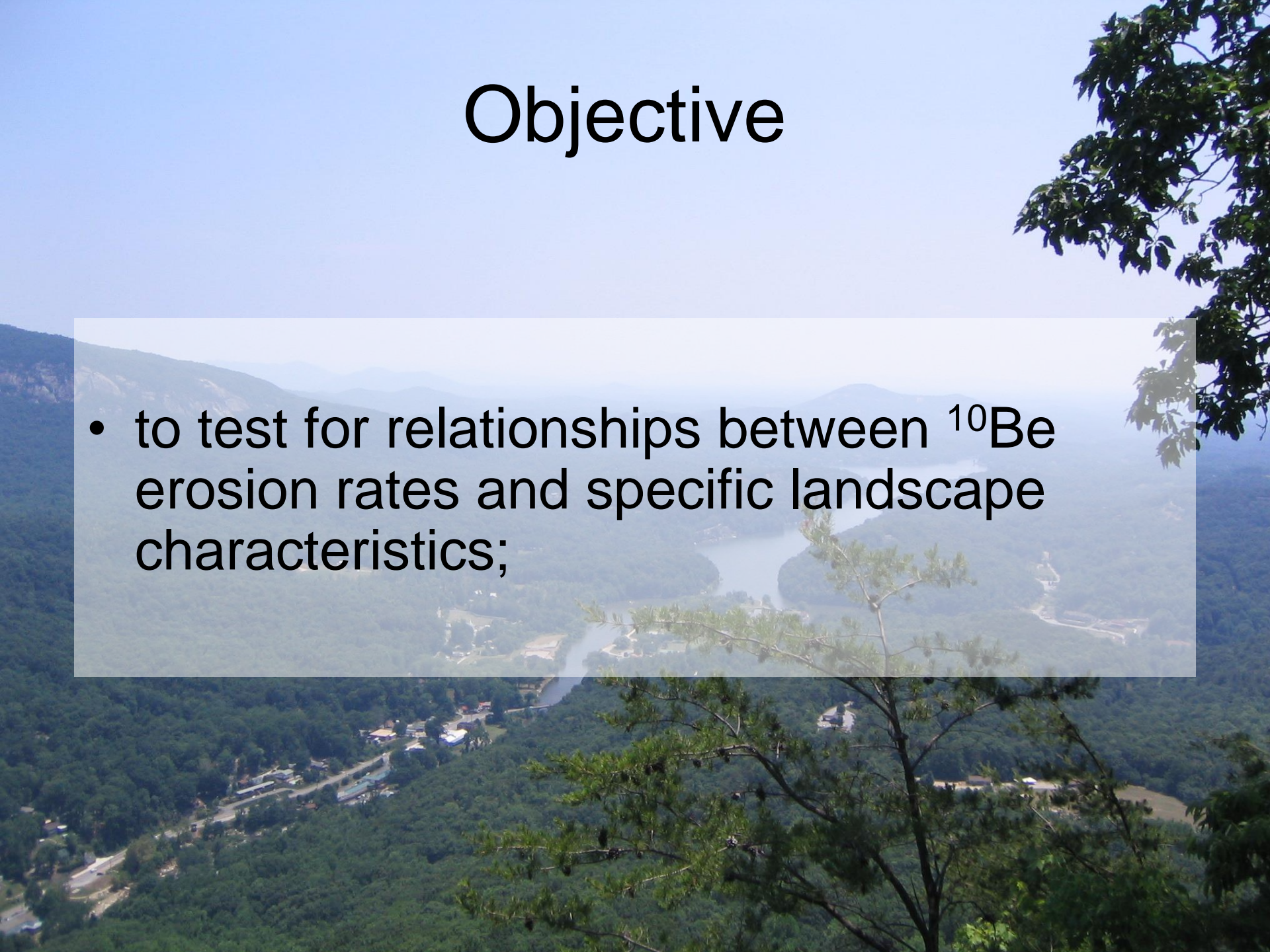
Blue Ridge (n=10)  
 $12.2 \pm 6.3 \text{ m My}^{-1}$

Piedmont (n=12)  
 $15.0 \pm 9.0 \text{ m My}^{-1}$

Erosion rates are slow! Consider that Wobus et al. (2005) measured  $180\text{-}770 \text{ m My}^{-1}$  in the central Nepalese Himalaya

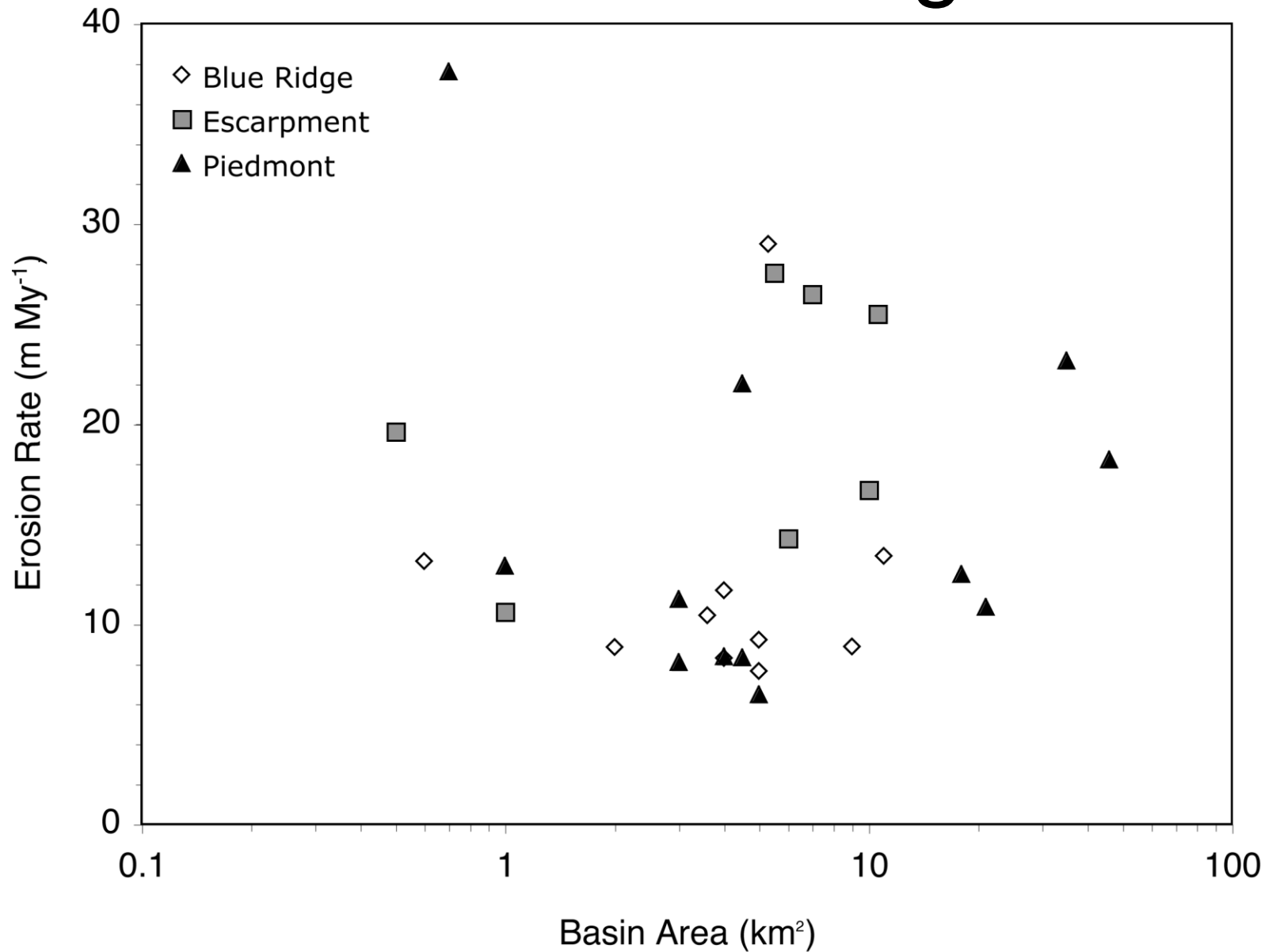
# Objective

- to test for relationships between  $^{10}\text{Be}$  erosion rates and specific landscape characteristics;

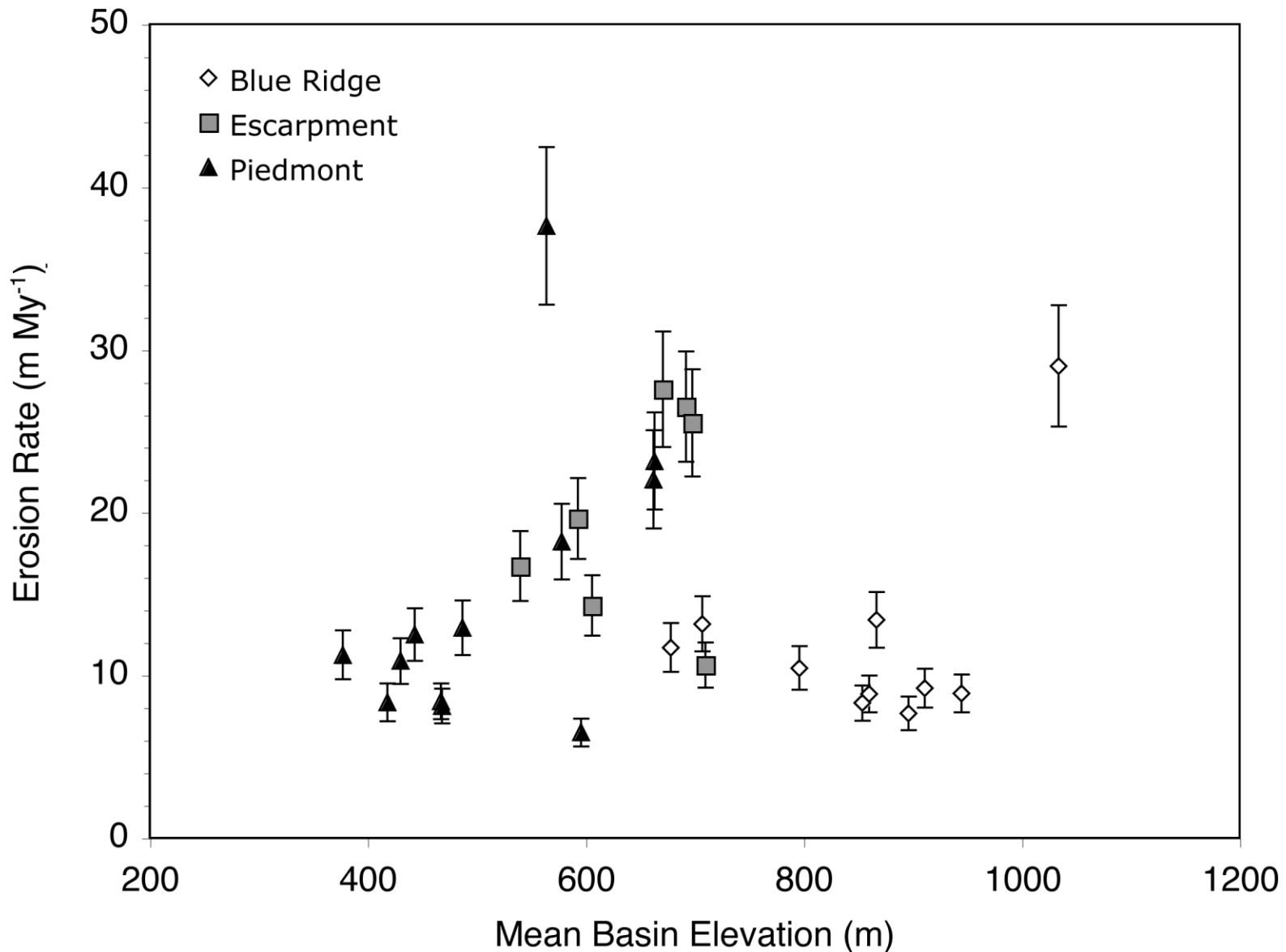




# Basin size is not significant

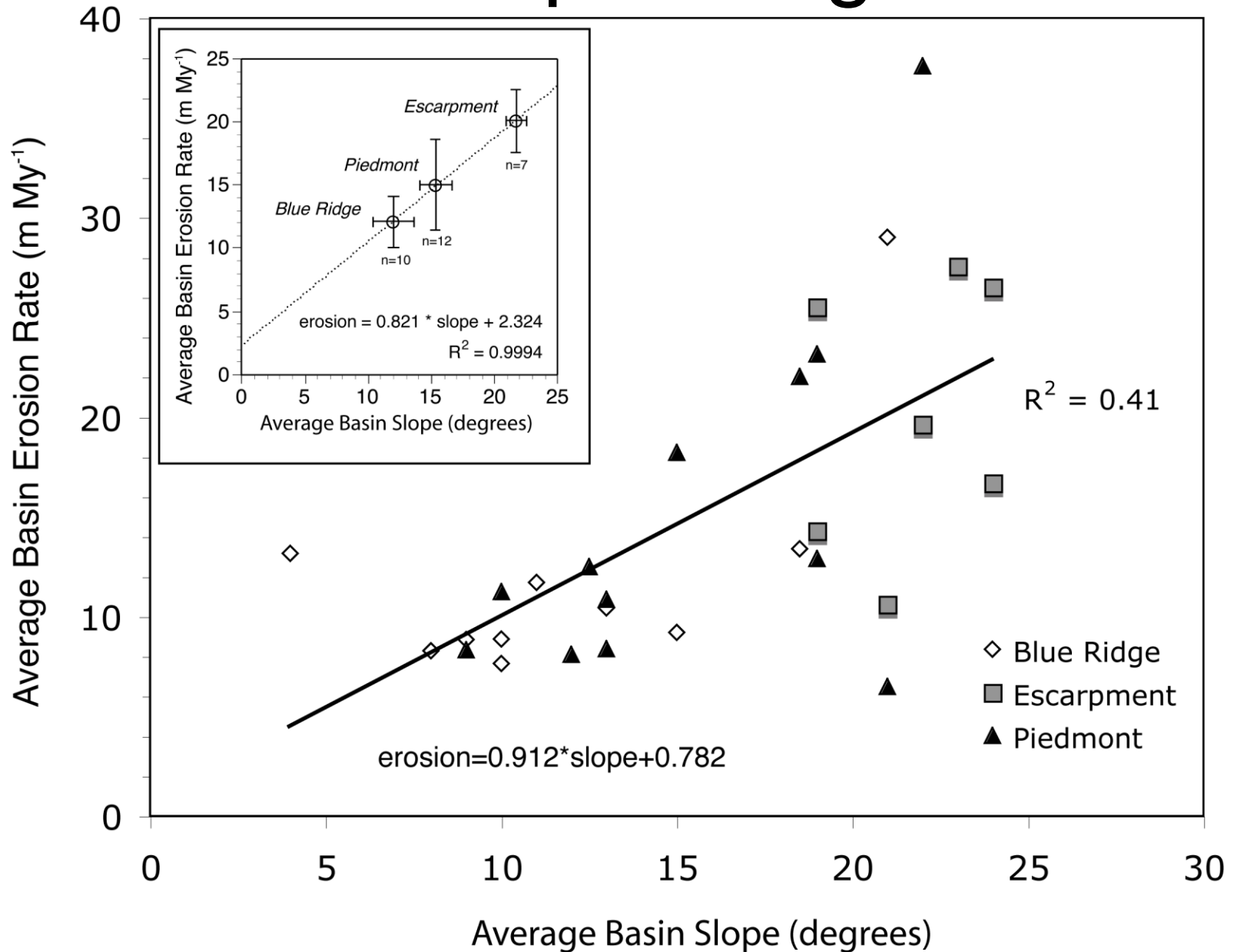


# Physiographic province shows slight relationship with erosion, not statistically separable





# Basin slope is significant



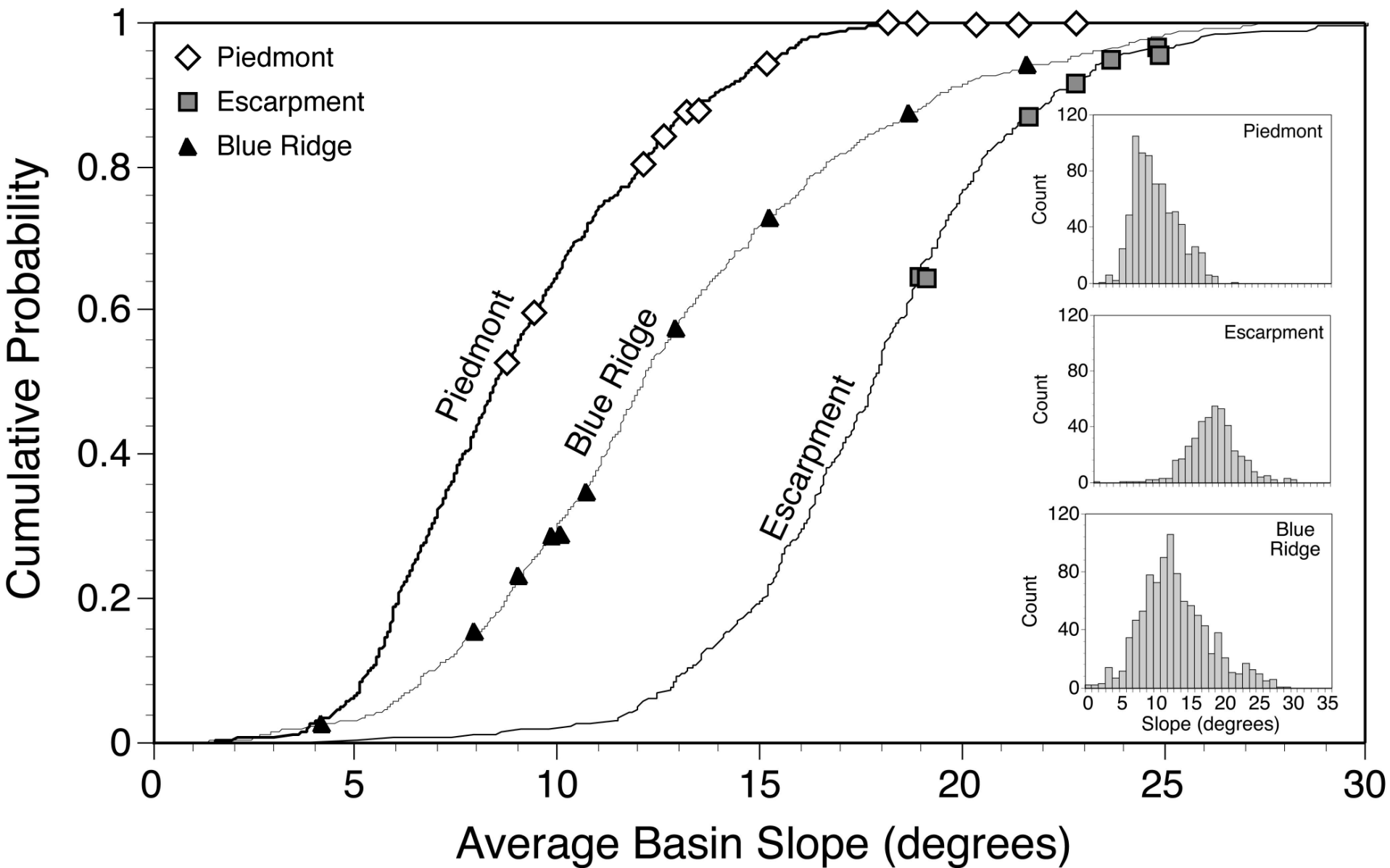
# Used GIS to assess average basin slope for the entire population of basins

	Population Slope	Sample Slope
Blue Ridge	12.8° n=968	12.0° n=10
Escarpment	17.7° n=428	21.7° n=7
Piedmont	9.0° n=738	15.3° n=12

Mean	5.6 km <sup>2</sup>	8.1 km <sup>2</sup>
Median	4.6 km <sup>2</sup>	5.0 km <sup>2</sup>



# Probability Density Function



# How can I model representative erosion rates?

- Erosion rate is dependant on slope:
  - Erosion rate = ( $^{\circ}$  slope) \* (0.912) + 0.78
- I calculated a model erosion rate for each province based on the slope of the average basin population.



# Integrated Model Erosion Rates

(n=7)  
 $20.1 \pm 6.6 \text{ m My}^{-1}$

Escarpment (n=428)  
 $17.1 \text{ m My}^{-1}$

Blue Ridge (n=968)  
 $12.5 \text{ m My}^{-1}$

(n=10)  
 $12.2 \pm 6.3 \text{ m My}^{-1}$

Piedmont (n=738)  
 $9.7 \text{ m My}^{-1}$

(n=12)  
 $15.0 \pm 9.0 \text{ m My}^{-1}$

# Objective

- to determine whether the BRE has evolved according to a model of ongoing & parallel retreat or by a model of rapid erosion following rifting and subsequent landscape stability.



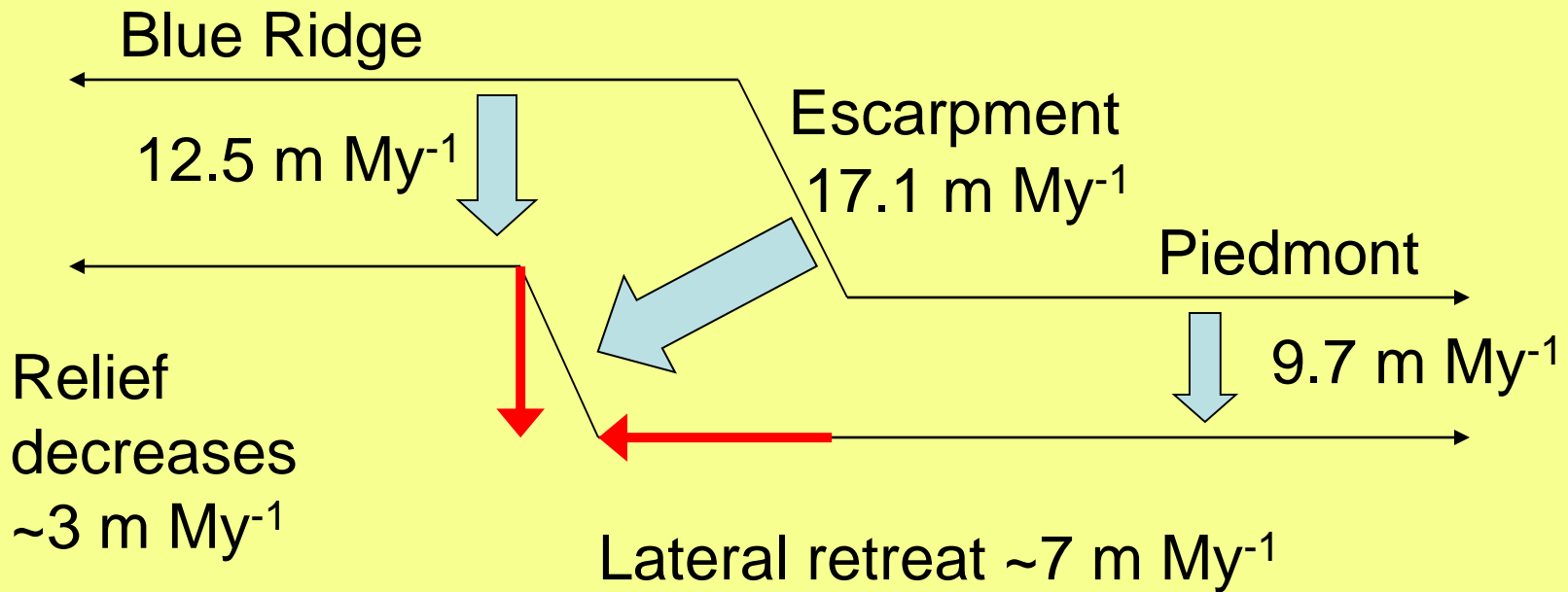
# Comparing Cosmogenic and Thermochronologic Erosion Rates

- $^{10}\text{Be}$  erosion rates are integrated over  $10^4$ - $10^5$  years
- Thermochronologic erosion rates are integrated over  $10^8$  years.
- Erosion Rate=**Depth** (integrated geothermal gradient and closure temp)/**Age** (U-Th)/He or # fission tracks.
  - AHe- (U-Th)/He closure temp 40-90°C
  - AFT- fission tracks closure temp 60-110°C

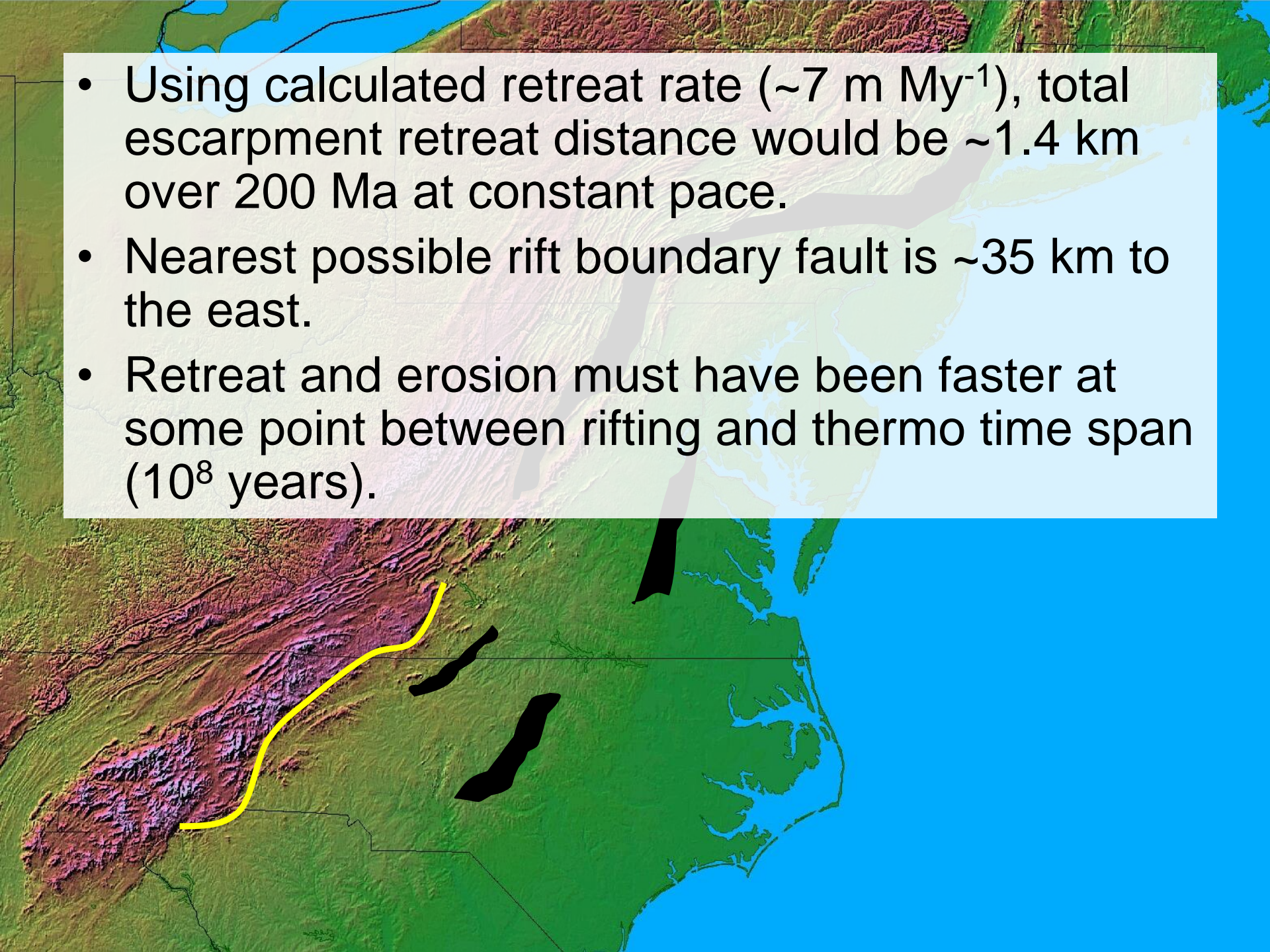
# Thermochronologic data consistent with Cosmogenic data

Both datasets are relatively slow





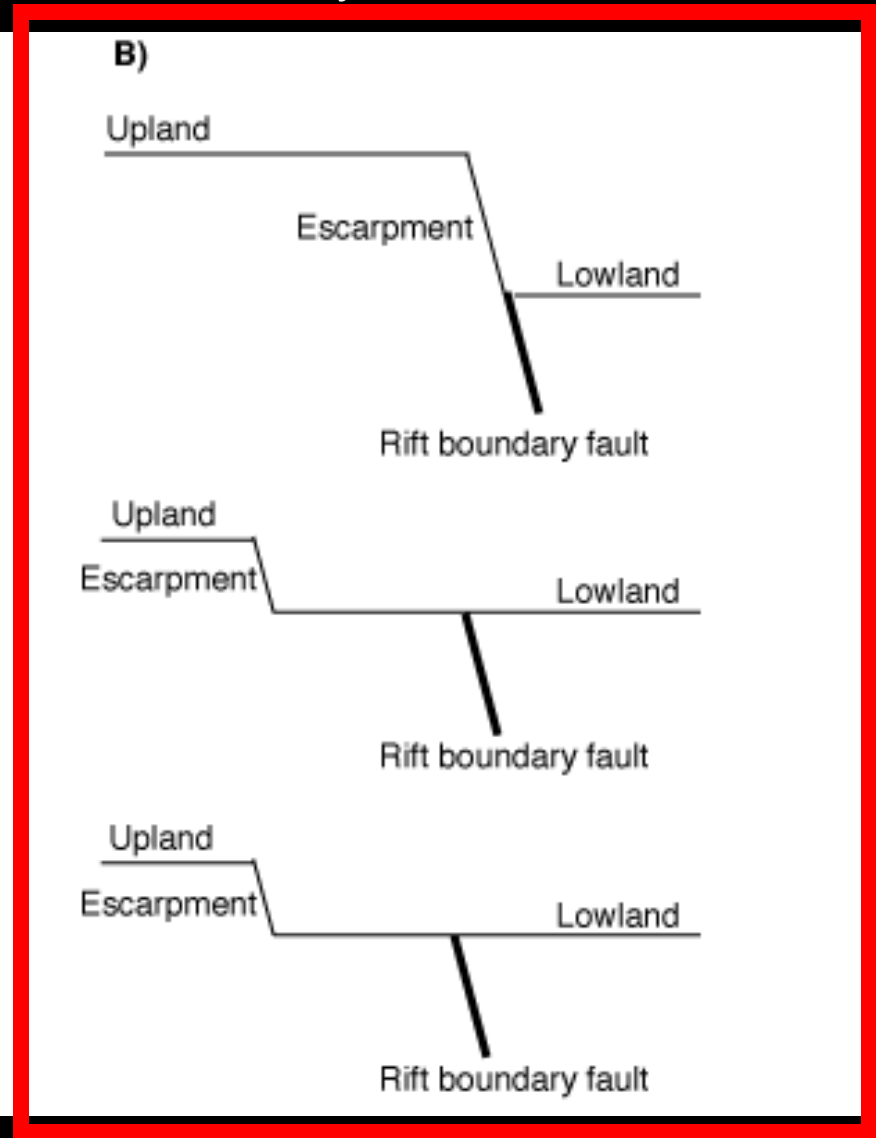
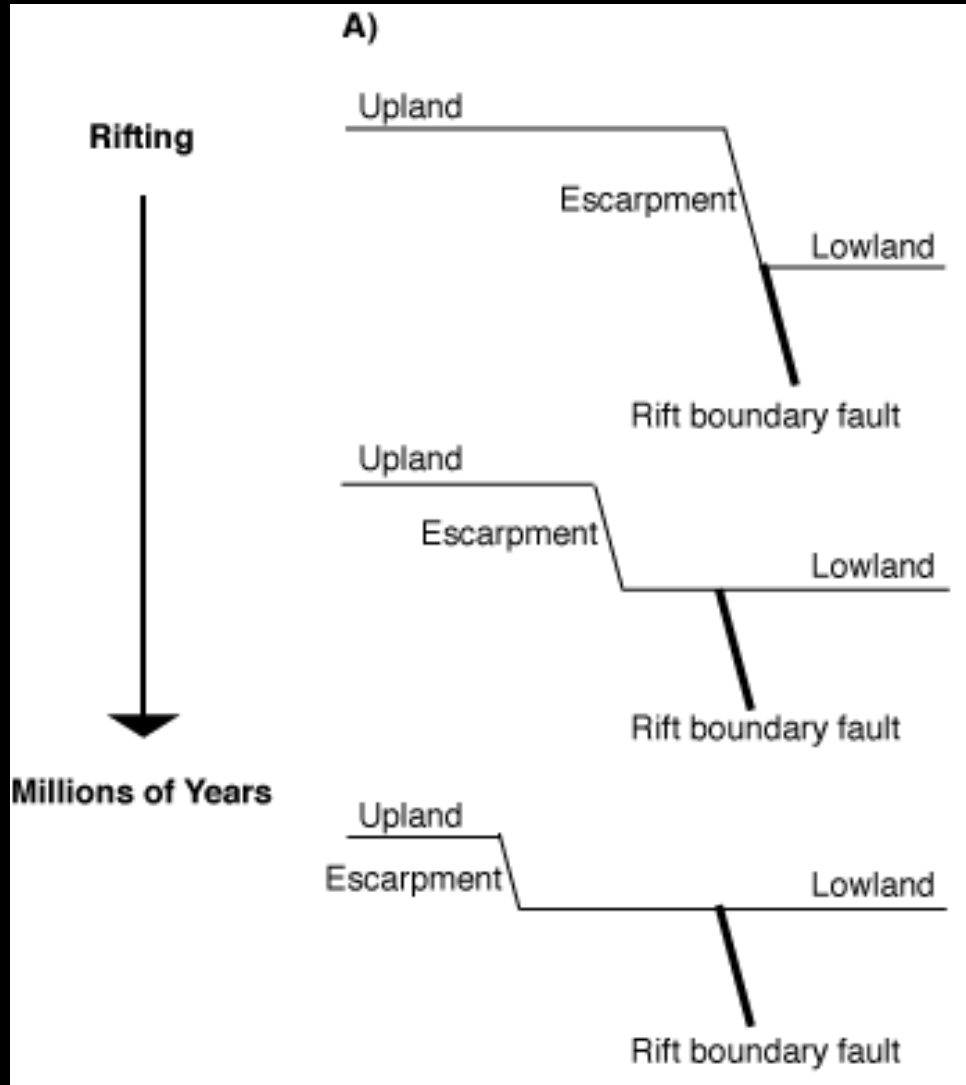
- Base level for the escarpment is set by the Piedmont therefore the difference in lowering rates can be taken as the retreat rate
- Piedmont is eroding more slowly than the Blue Ridge therefore relief is decreasing

- 
- A topographic map of the Appalachian region, showing elevation in shades of green, brown, and red. A yellow line traces a path from the western part of the range towards the east. Several black shapes are overlaid on the map, representing geological features or faults. The map is partially obscured by a white text box in the upper left.
- Using calculated retreat rate ( $\sim 7 \text{ m My}^{-1}$ ), total escarpment retreat distance would be  $\sim 1.4 \text{ km}$  over 200 Ma at constant pace.
  - Nearest possible rift boundary fault is  $\sim 35 \text{ km}$  to the east.
  - Retreat and erosion must have been faster at some point between rifting and thermo time span ( $10^8$  years).



Ongoing & steady retreat

Significant retreat following rifting, then stability



# Other escarpments

- These results agree with studies from other passive margin escarpments such as:
  - **Namibia** (Bierman and Caffee, 2001; Brown et al., 2000)
  - **South Africa** (Fleming et al., 1999; Summerfield et al., 1997)
  - **Southeastern Australia** (Heimsath et al., 2006; Persano et al., 2002)
  - **Sri Lanka** (Vanacker et al., 2007)



# Conclusions

- Grain size does not affect  $^{10}\text{Be}$  concentration on and near the BRE
- Overall the BRE is lowering and retreating very slowly
- Average slope is the only basin characteristic that influences erosion on and near the BRE
- The BRE appears to have evolved through a period of significant and rapid erosion immediately following rifting and has remained a fairly stable feature of the landscape since that time.



# Acknowledgments

- NSF
- Paul Bierman
- Milan Pavich & Scott Southworth (USGS)
- Donna Rizzo & Keith Klepeis
- Jen Larsen, Luke Reusser, Jane Duxbury & Matt Jungers
- Corey Coutu
- Dave Linari







# Rocks Older Than Dinosaurs?









# Cosmogenic isotope production with depth

$$P_x = P_0 e^{-(x\rho/\Lambda)}$$

Variables:

$P_x$  = nuclide production rate at depth  $x$

$P_0$  = sediment production rate (5.17 atoms  $\text{g}^{-1} \text{y}^{-1}$ )

$\rho$  = density of material (2.7  $\text{g cm}^{-3}$  for rock)

$\Lambda$  = attenuation factor (165  $\text{g cm}^{-2}$ )

# Erosion rate calcs...

$$m/\rho = \varepsilon = \Lambda(P - N)/\rho N$$

Variables:

$\varepsilon$  = erosion rate

P = basin effective production rate

$\rho$  = density of material (2.7 g cm<sup>-3</sup> for rock)

$\Lambda$  = attenuation factor (165 g cm<sup>-2</sup>)